

**Current issues in Atmospheric Chemistry and Climate**  
**Aktuelle Problemstellungen in Atmosphärenchemie/physik in Bezug**  
**auf den Klimawandel**



***Thomas Karl***

***Institute for Meteorology and Geophysics – University of Innsbruck***

# Atmosphere



**Diameter of the atmosphere: 1999 km (p = 1 atm)**  
**Mass: 5140 Trillion tons**

*PhotoCredit: NASA*

# Atmosphere



**Volume: approx. 3% compared to planet ( $p = 1$  atm)**  
**Mass: only 0.00009 % of the Earth's mass**

# Atmosphere



due to its small reservoir the atmosphere is the most vulnerable part of the planet – any environmental changes on the planet manifest themselves fastest in the atmosphere

# NMVOOC

Definition:

NMVOOC: non methane volatile organic compounds (often also termed VOC)

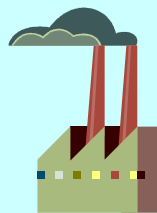
Other subcategories used:

- BVOC: biogenic volatile organic compounds
- oVOC: oxygenated volatile organic compounds
- NMHC: non methane hydrocarbons (old terminology)



## NMVOOC

Anthropogenic



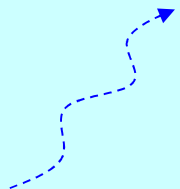
Biomass burning



Biogenic/natural



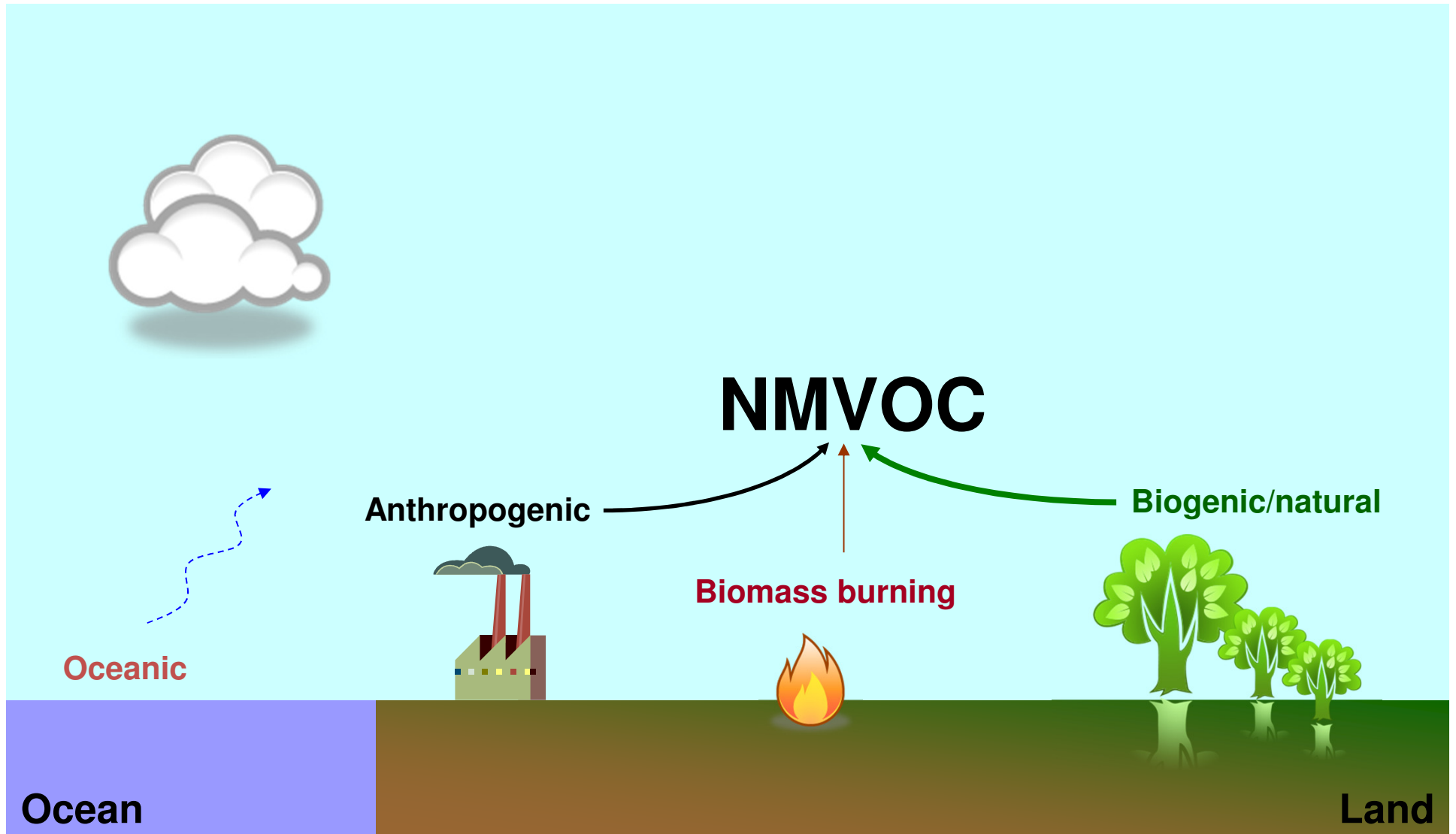
Oceanic



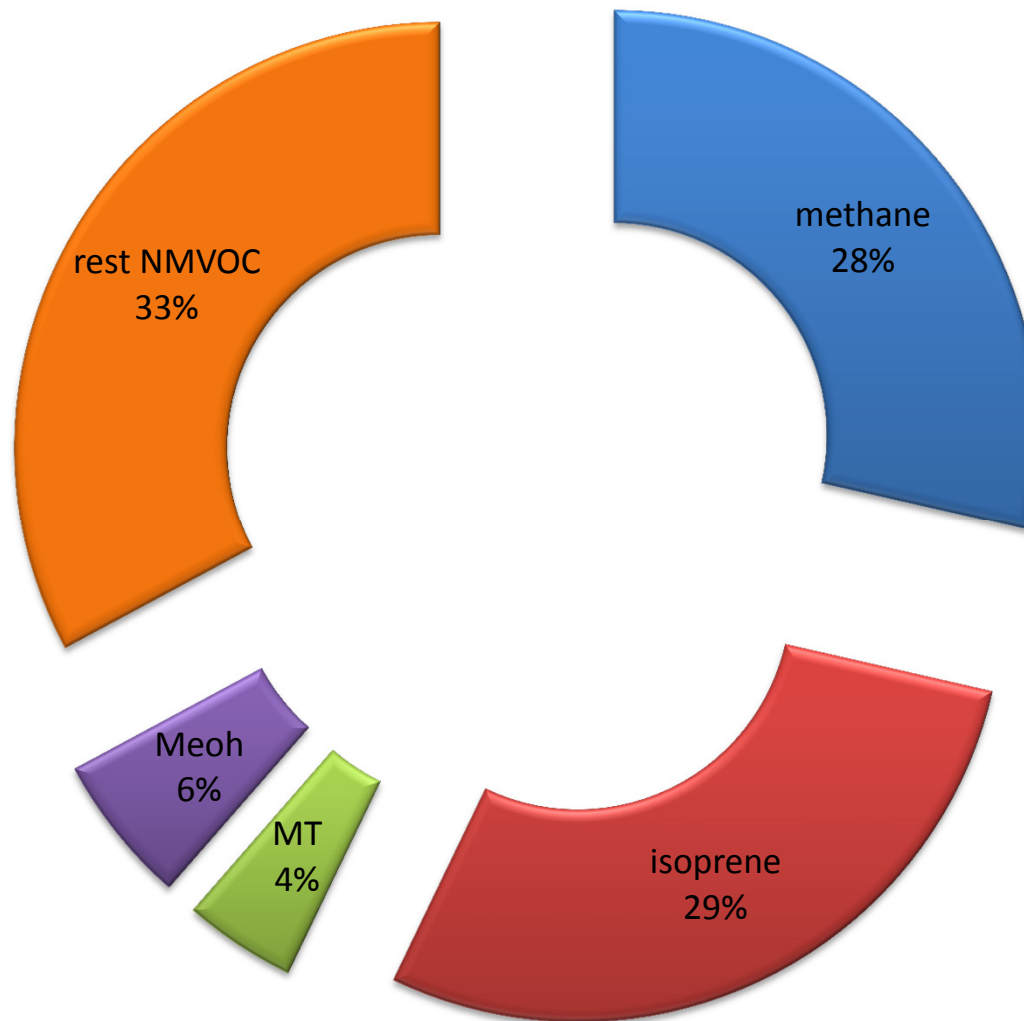
Ocean

Land

# Budgets of NMVOC



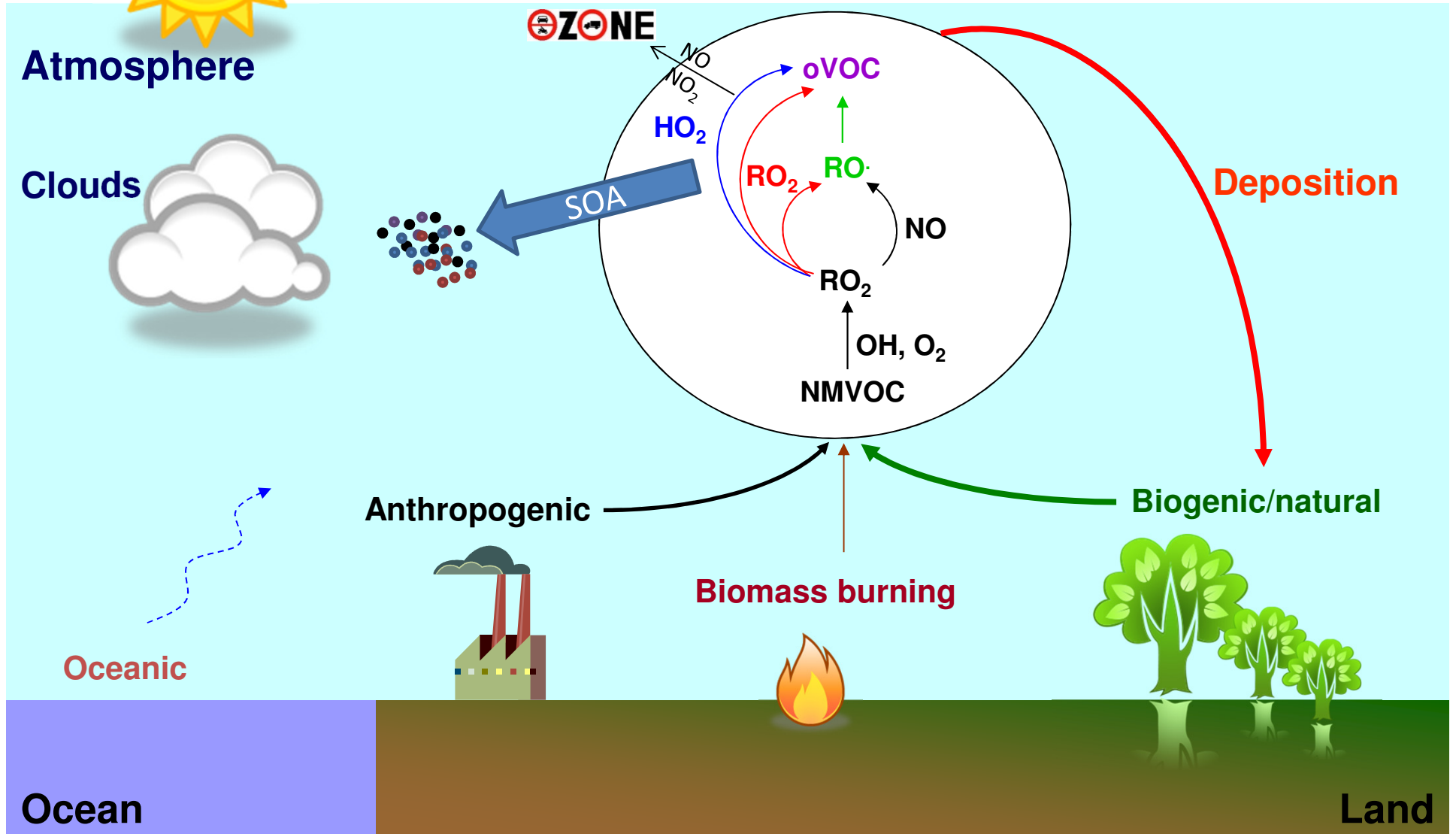
# Budgets of Methane and NMVOC Fluxes



# Earth System

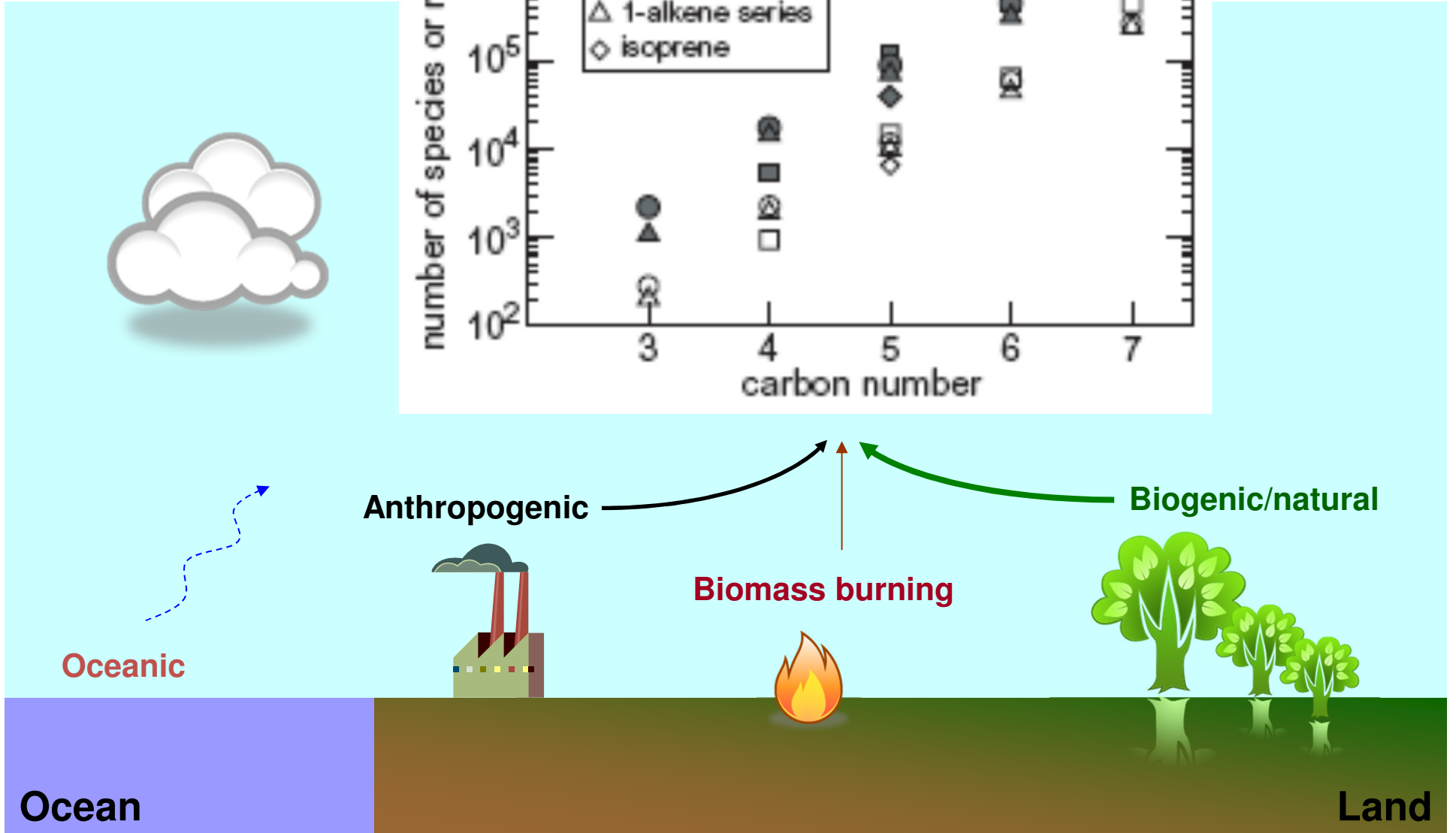
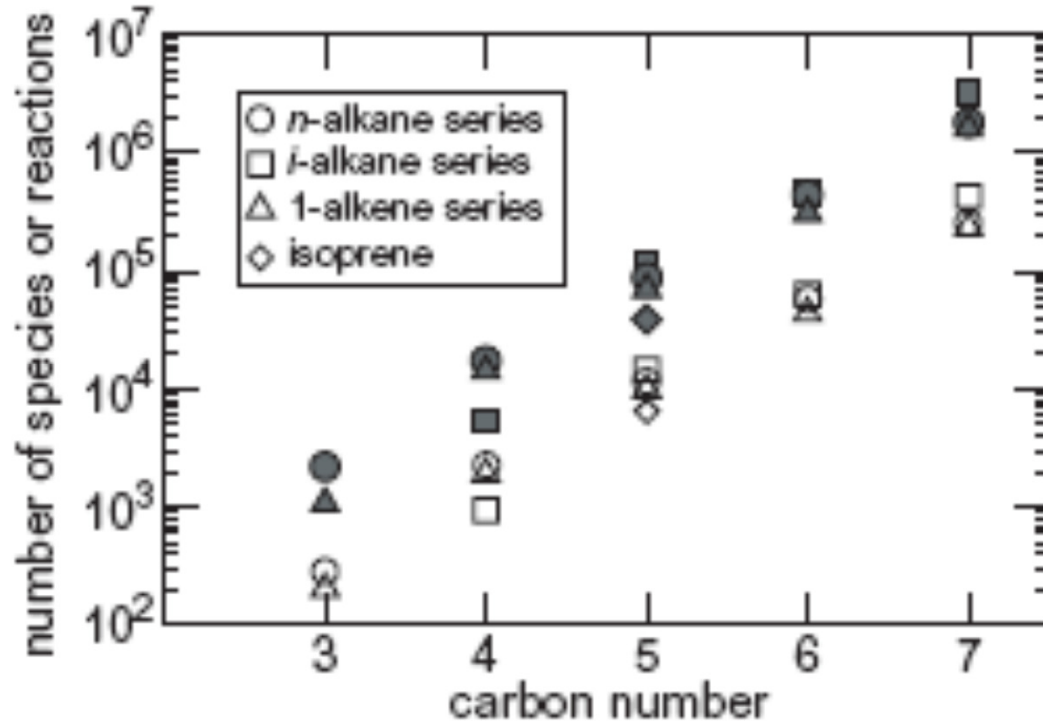


SOA: secondary organic aerosol





Aumont et al., ACP, 2005

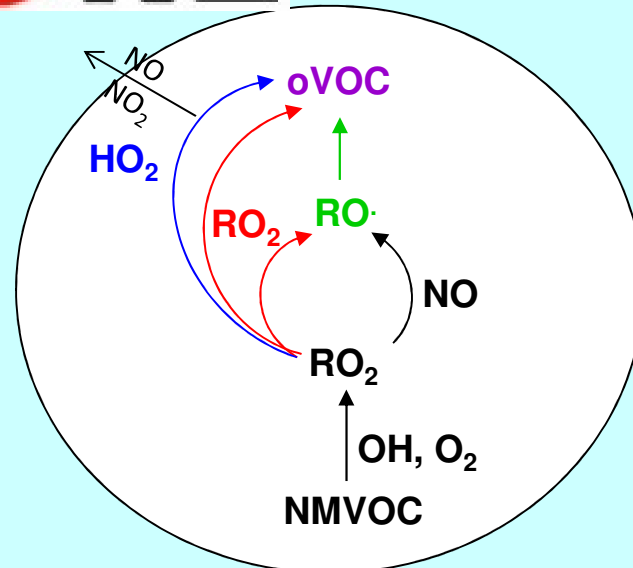


# Earth System



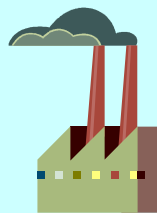
# **OZONE**

Atmosphere



Oceanic

Anthropogenic



Biomass burning



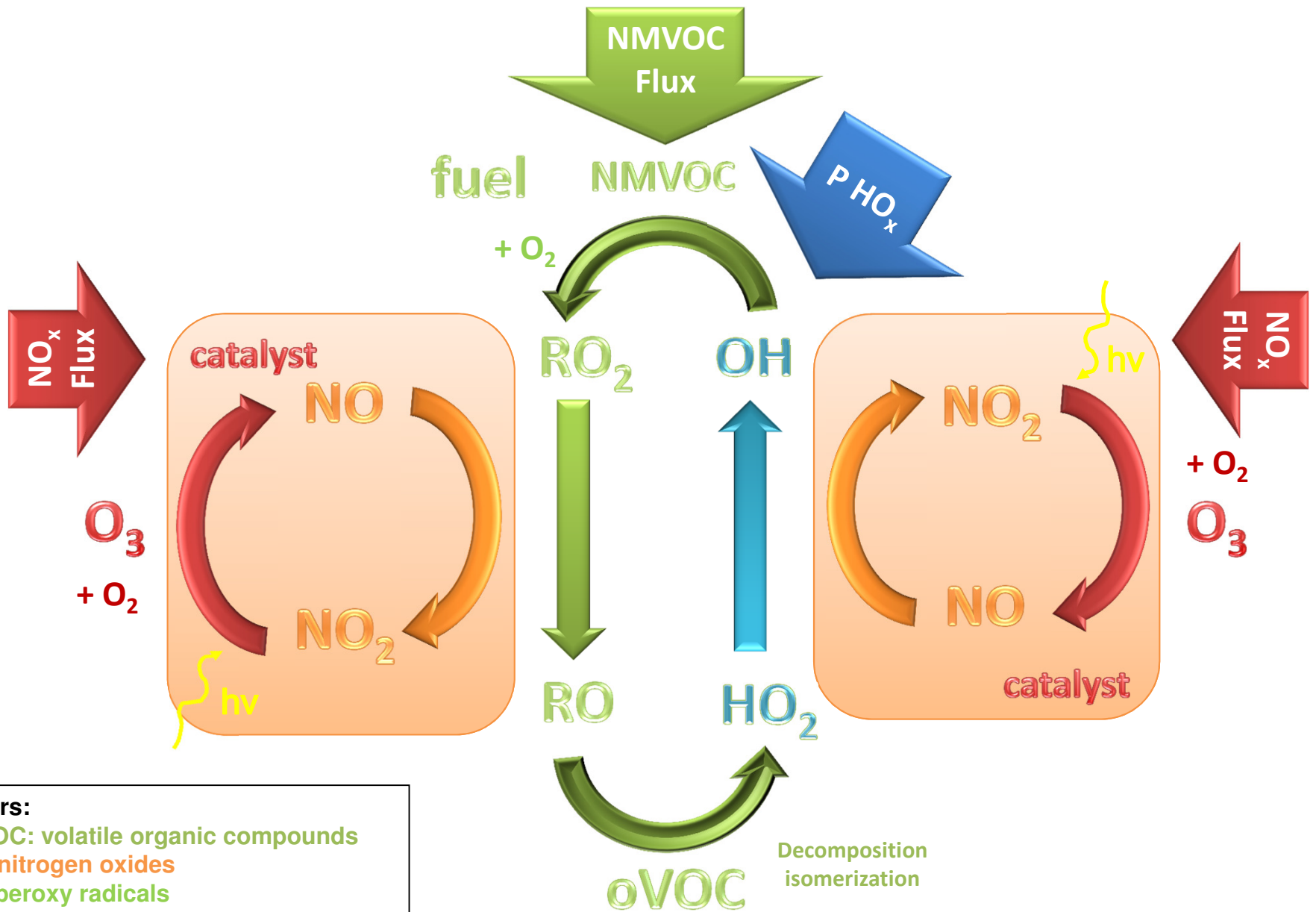
Biogenic/natural



Ocean

Land

# The photochemical engine



**Players:**  
 NMVOC: volatile organic compounds  
 NO<sub>x</sub>: nitrogen oxides  
 RO<sub>x</sub>: peroxy radicals  
 HO<sub>x</sub>: hydroxy radicals  
 hv: radiation

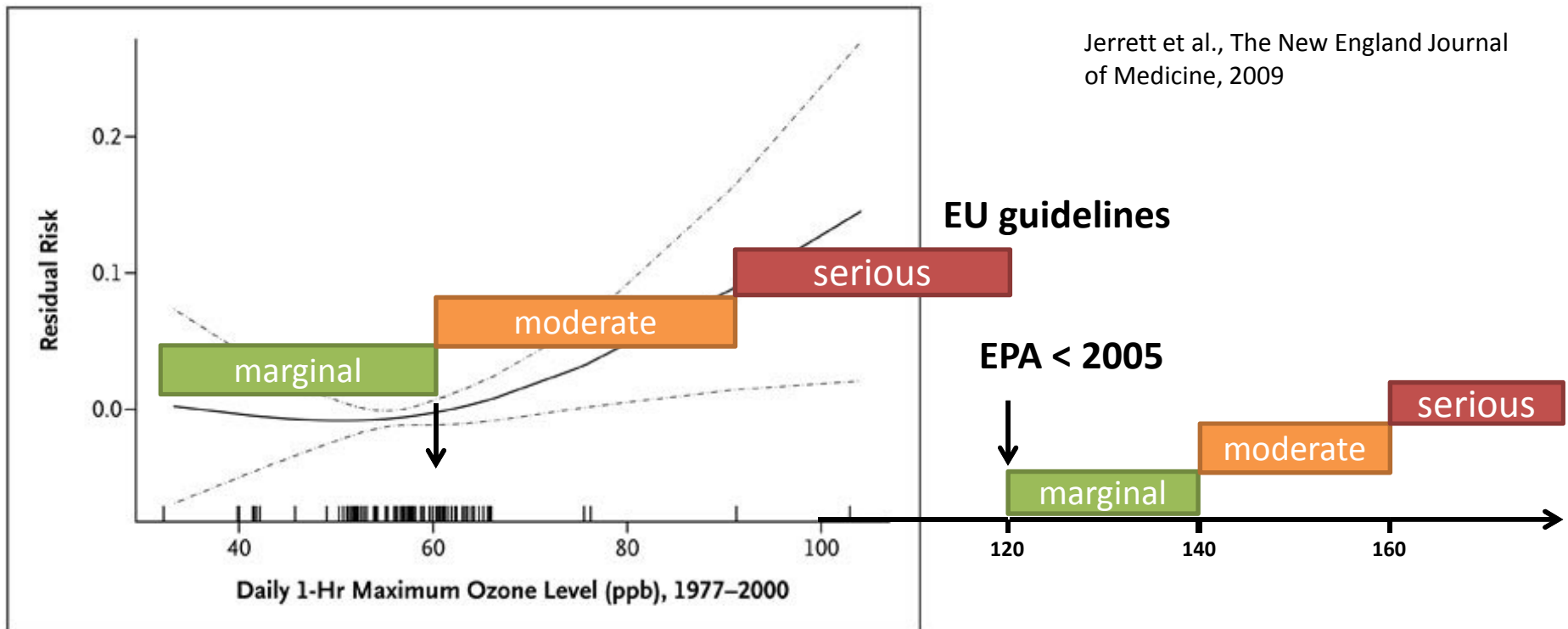


## Elevated Ozone Causes Health Problems and Damages Plants/Crops

NMVOCs fuel an oxidizing atmosphere

e.g. in US: ~14-55 billion US\$ health care costs / year

### Long –Term Ozone Exposure and Mortality .

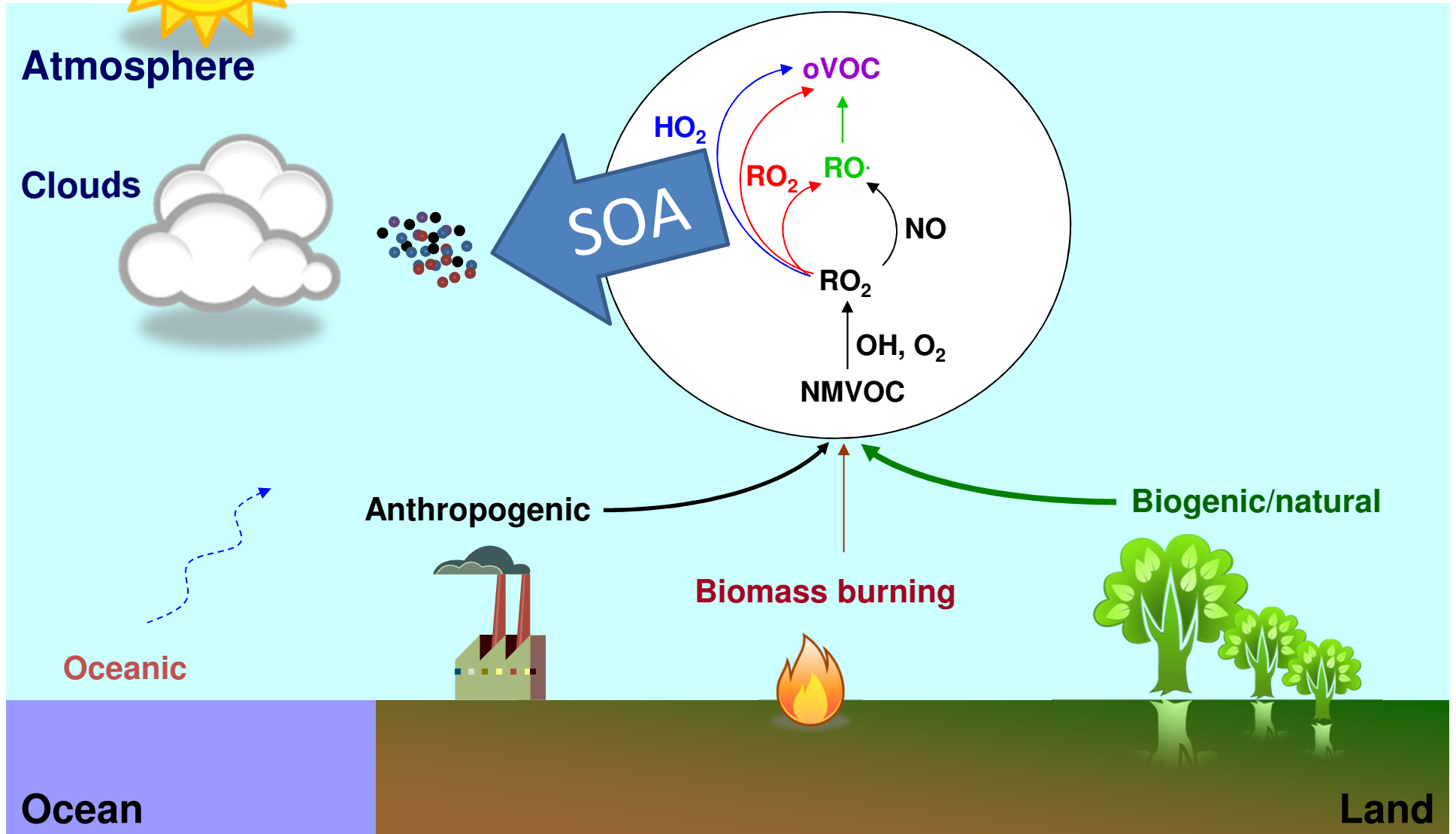


# Earth System



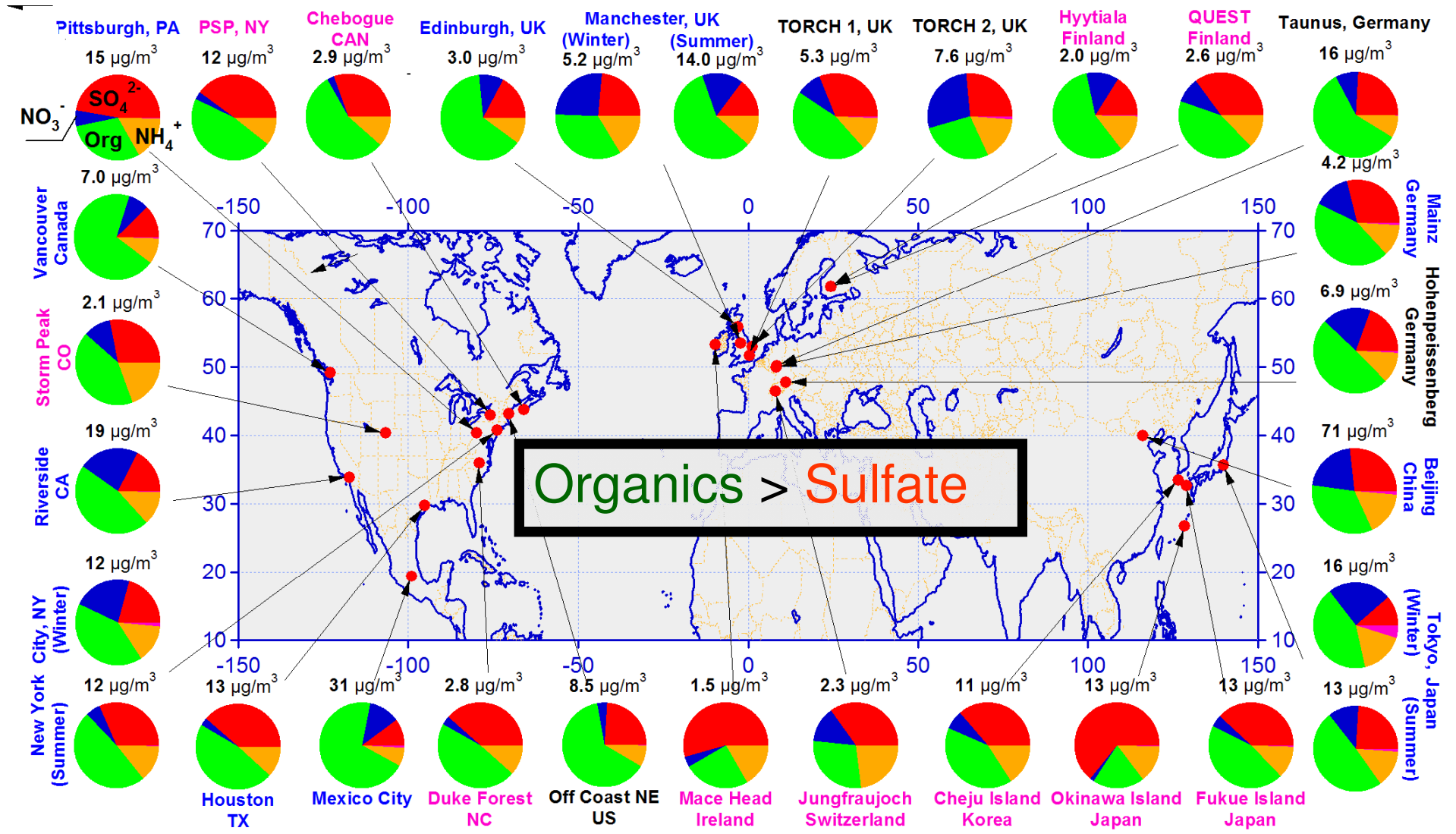
# Aerosols

Secondary organic aerosol



# Atmospheric oxidation of NMVOCs leads to organic aerosol formation

Accumulation Mode  $\approx 1 \mu\text{m}$



Zhang, Jimenez et al., *GRL*, 2007

# Dominant aerosol sources

- Dust
- Sea Salt
- Organic Carbon (POA+SOA)
- Black carbon (BC)
- Sulfur

POA: primary organic aerosol

SOA: secondary organic aerosol

Source *	Total (Tg/y)
Dust	1000-3000
Sea salt	8000-16000
SOA	100-1500?
POA	~1000 <sup>x</sup>
BC	6-8
Sulfur	16-45

\* IPCC 4/5

<sup>x</sup> Jaenicke et al., Science, 2005

# Dominant aerosol sources

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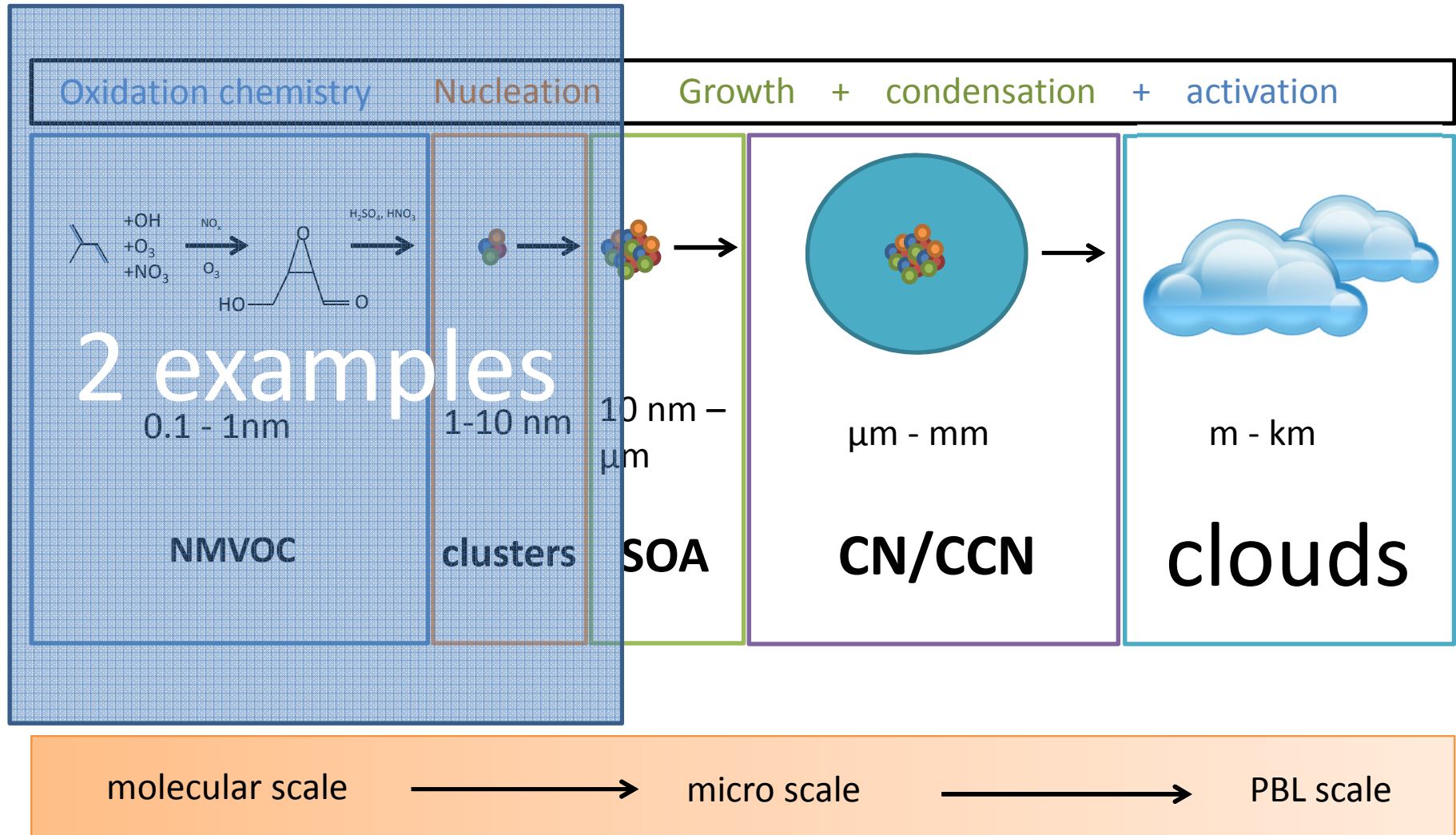
Source*	Total (Tg/y)	Accumulation mode <2.5µm (Tg/y)
Dust	1000-3000	70-700 (270)
Sea salt	8000-16000	1200-2400 ??
<b>SOA</b>	<b>100-1500?</b>	<b>100-1500 (850??)</b>
POA	~1000 <sup>x</sup>	<<
BC	6-8	<6 (?)
Sulfur	16-45	16-45

\* IPCC 4/5

<sup>x</sup> Jaenicke et al., Science, 2005



# From molecules to clouds



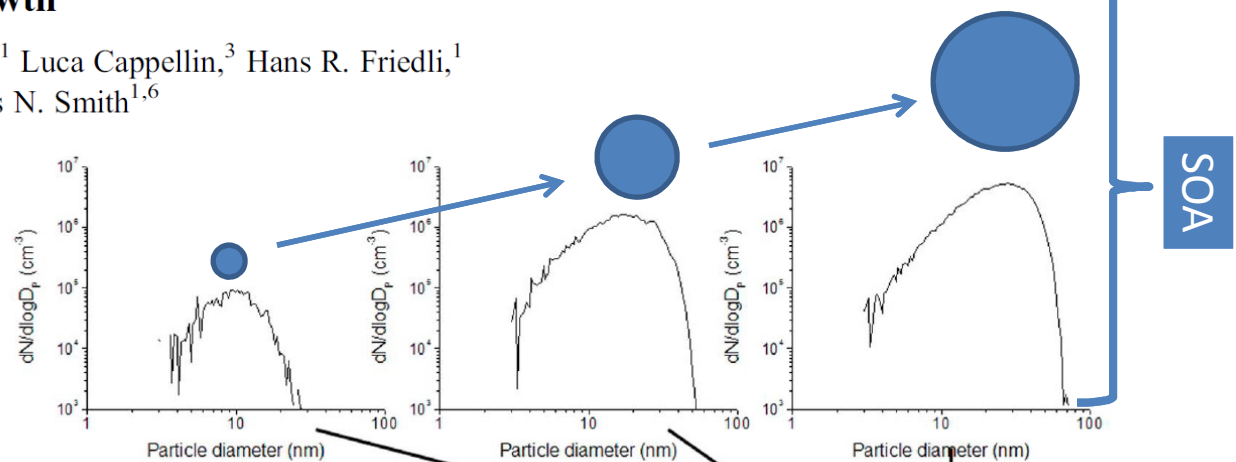
# Example 1 of biogenically enhanced SOA in ultrafine mode

## Identification of the biogenic compounds responsible for size-dependent nanoparticle growth

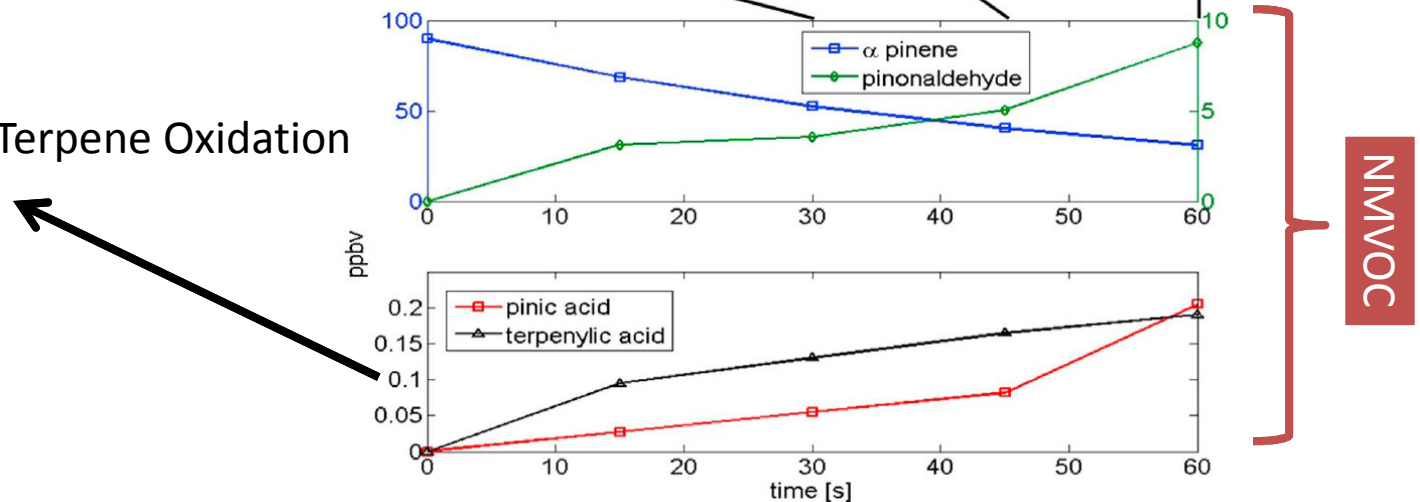
Paul M. Winkler,<sup>1,2</sup> John Ortega,<sup>1</sup> Thomas Karl,<sup>1</sup> Luca Cappellin,<sup>3</sup> Hans R. Friedli,<sup>1</sup> Kelley Barsanti,<sup>4</sup> Peter H. McMurry,<sup>5</sup> and James N. Smith<sup>1,6</sup>

GRL, 2012

Evolution of particle size



Organic acids from Terpene Oxidation



## Example 2: Studying the influence of GCR (galactic cosmic rays) on aerosol production and cloud formation at CERN

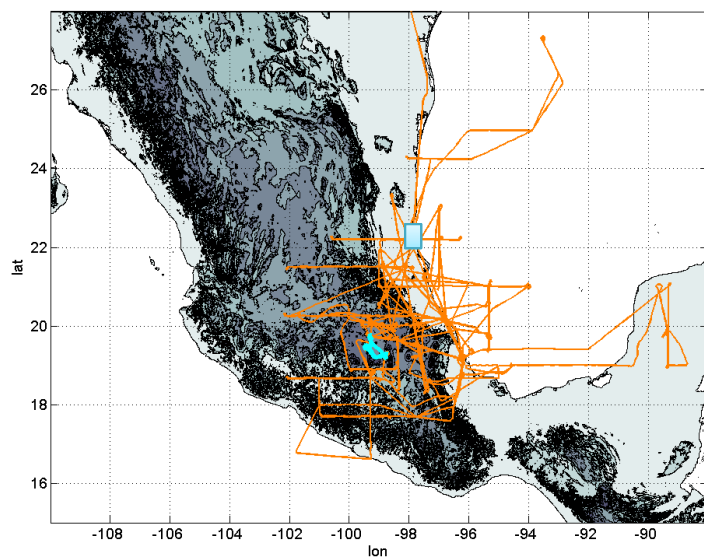


The CLOUD project at CERN

<http://cloud.web.cern.ch/cloud/Physics/modulation.html>

Results: GCR can not explain the formation rate of ultrafine aerosols in the lower atmosphere!  
NEED precursor gases

# Brown Haze over Mexico City



MIRAGE, 2006

# Blue Haze



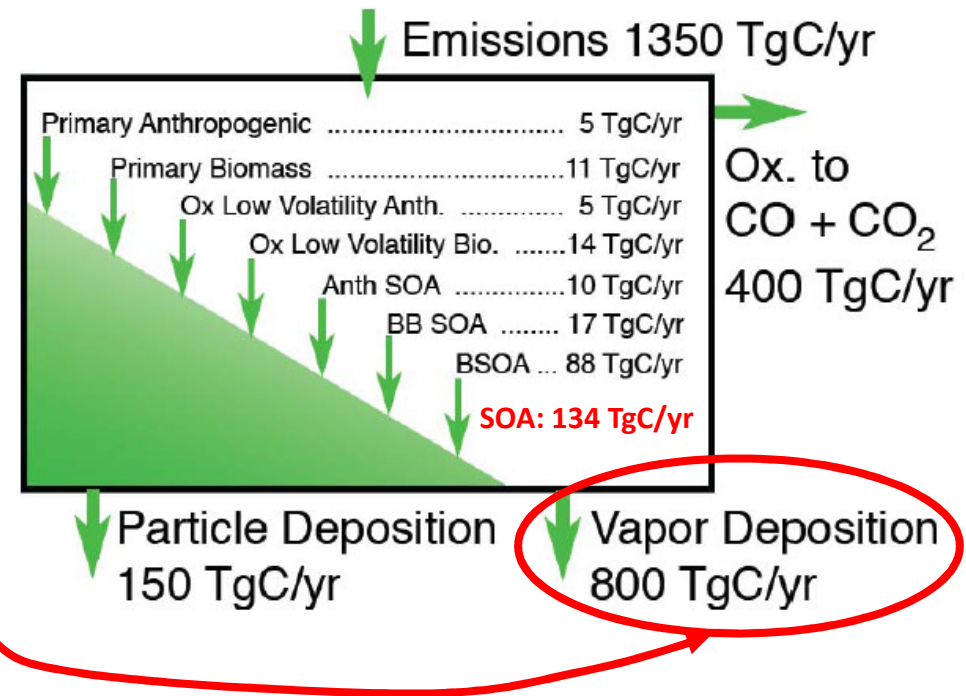
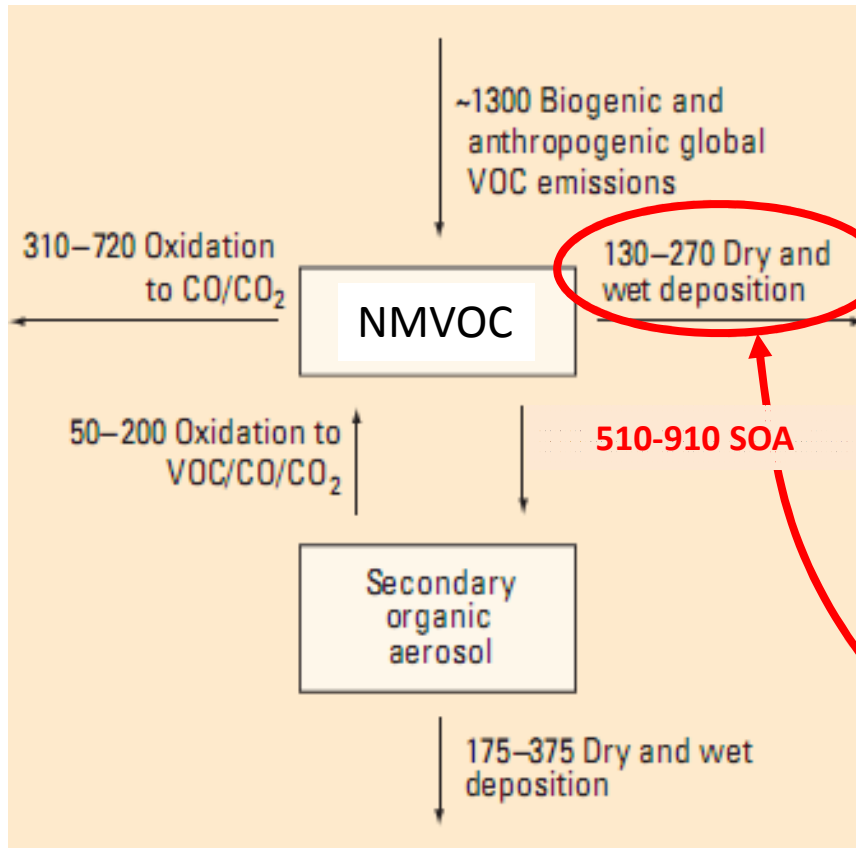
*Blue Haze above the OZARKS*

Ozarks, USA, 2013  
Thomas Karl

# Two different estimates of SOA: 710 TgC/y vs 134 TgC/y

Goldstein and Galbally, ES&T, 2007

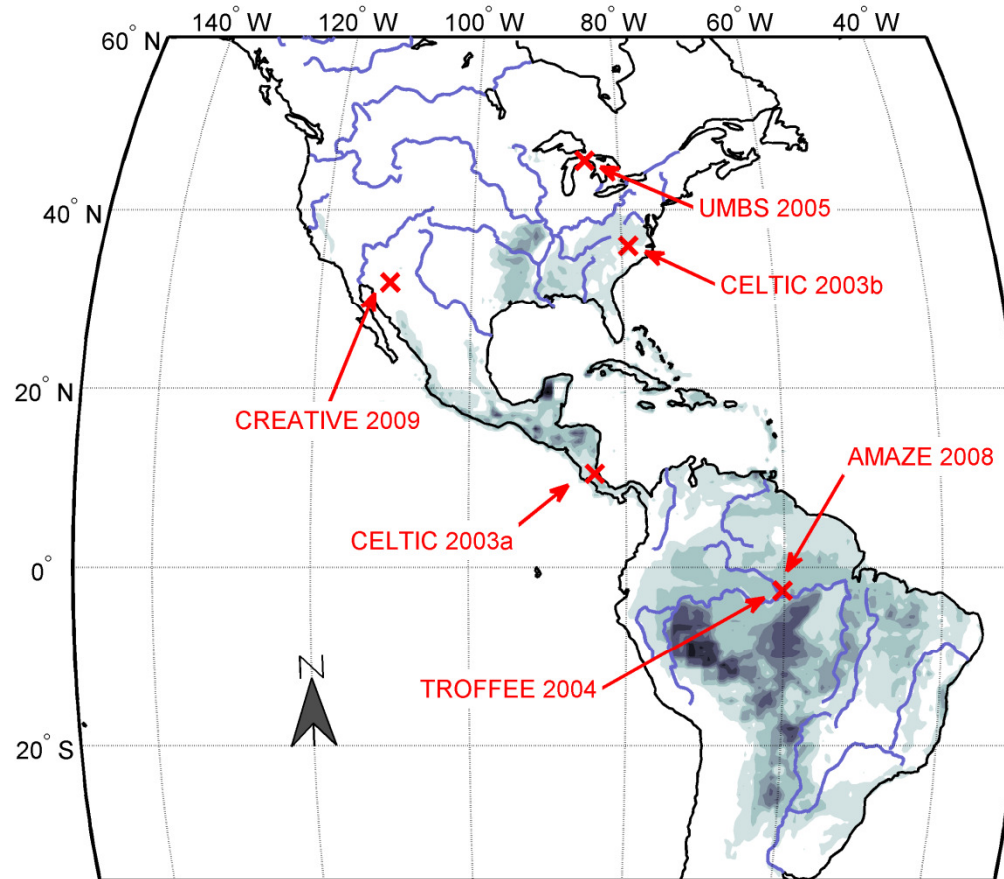
Hallquist et al., ACP, 2009



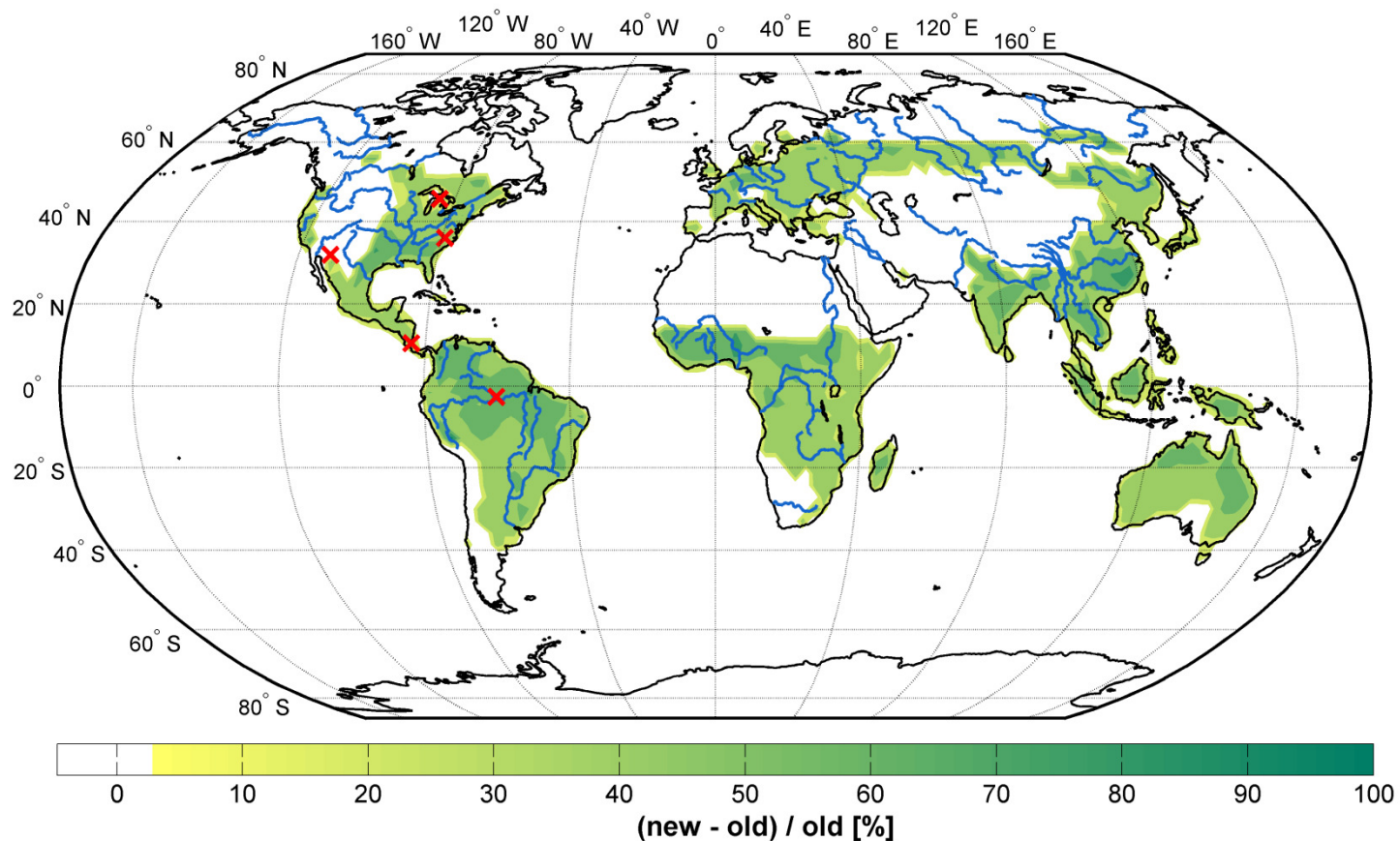
**large uncertainty in deposition assumptions**



# *Field campaigns 2003-2009*







## 1300 TgC/a primary VOC input

	Mean [TgC/a]	Comments
<b>this study</b>	<b>590±130</b>	<b>Dry and wet deposition (vapors)</b>
Goldstein and Galbally (2007)	200±100	Dry and wet deposition (vapors)
Hallquist et al., 2009	800	Dry and wet deposition (vapors)
Willey et al. (2000)	430±150	Wet deposition (vapors+particles)

# US Environmental Protection Agency

Figure 3. Rate of Temperature Change in the United States, 1901–2012

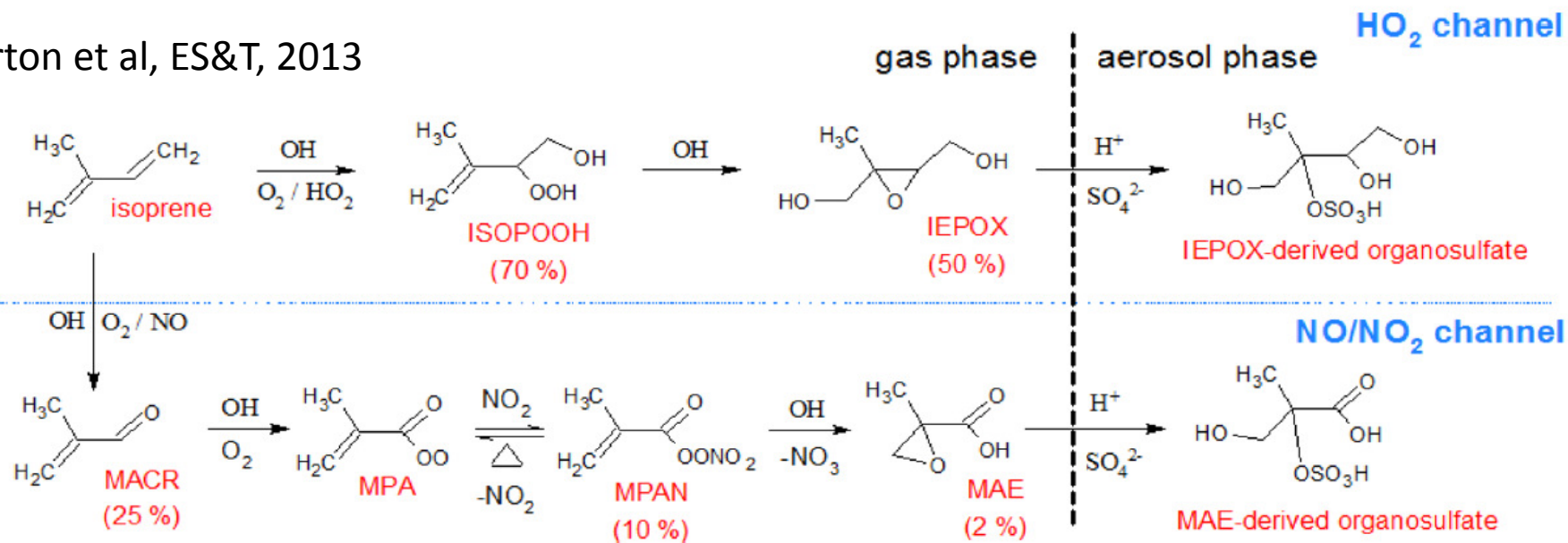


warmed like S.

# Blue Haze formation via terpene derived organosulfates?

Need anthropogenic pollution (e.g. SO<sub>2</sub>, or NO<sub>x</sub>), which interacts and enhances the formation of natural aerosol formation

Worton et al, ES&T, 2013

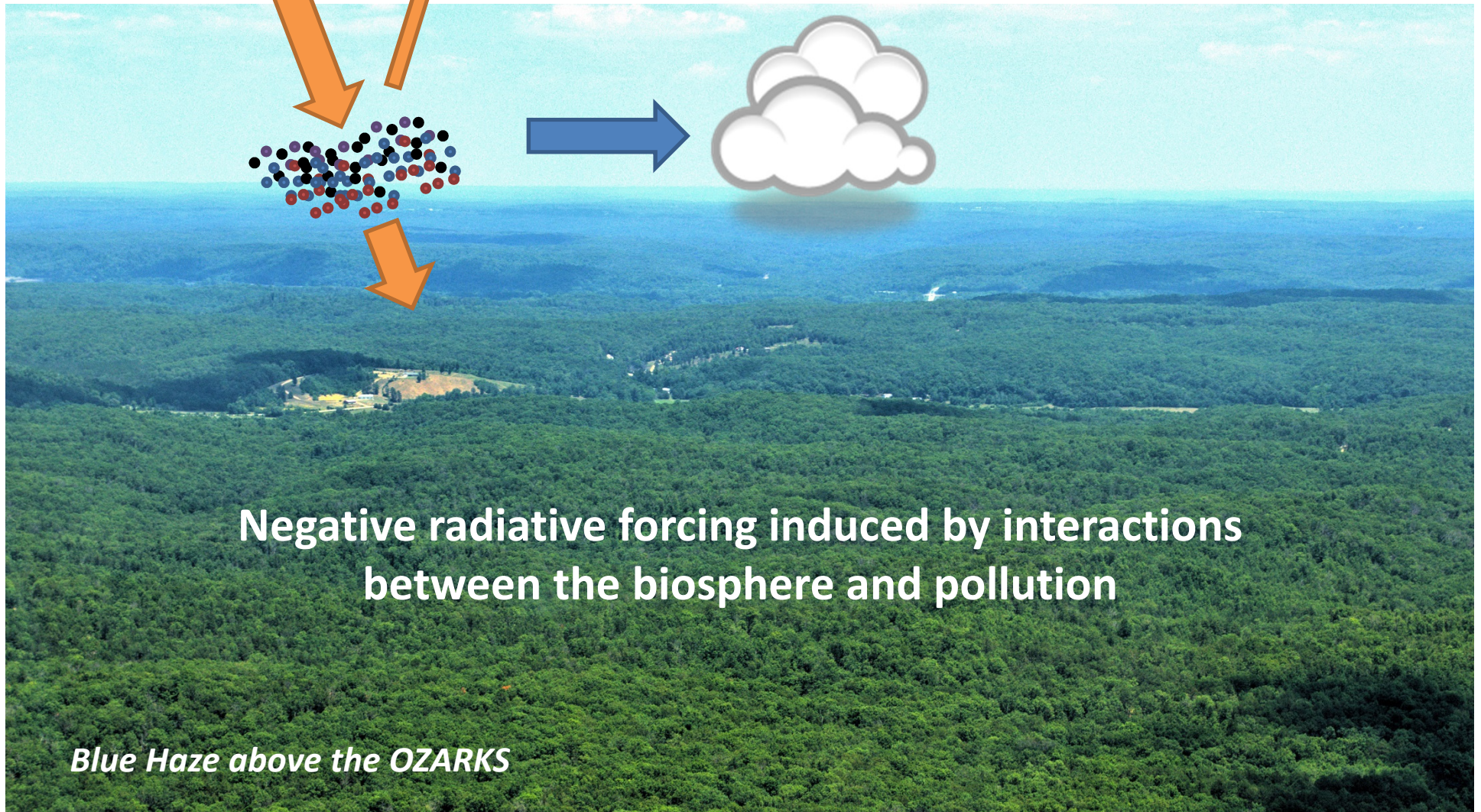


# Blue Haze formation via terpene derived organosulfates?



*Blue Haze above the OZARKS*

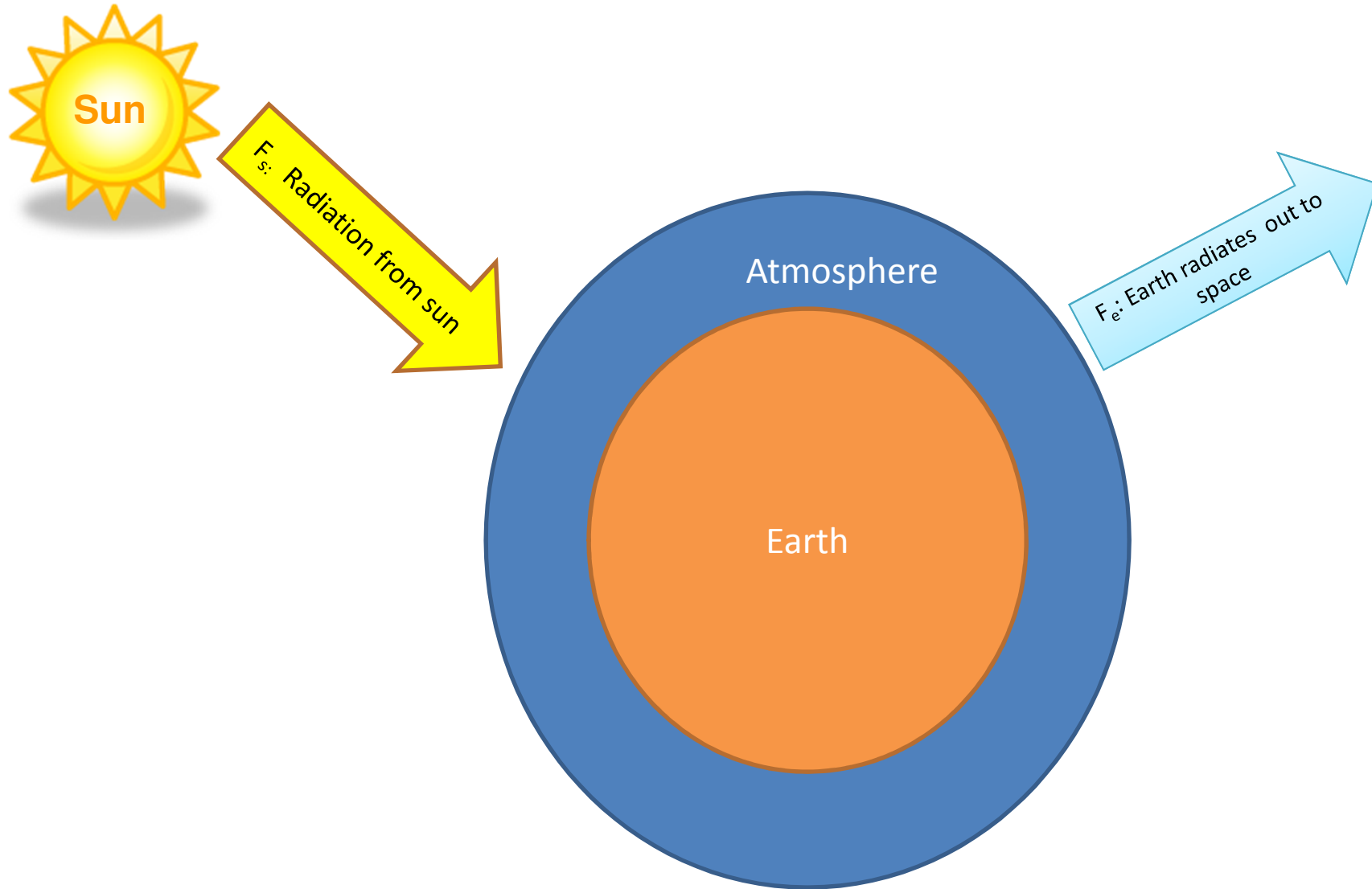
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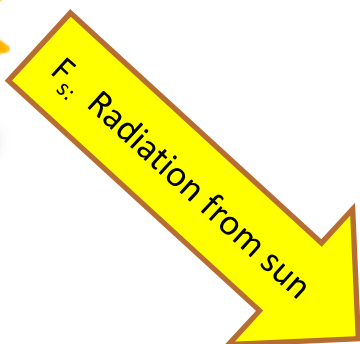


Negative radiative forcing induced by interactions between the biosphere and pollution

*Blue Haze above the OZARKS*

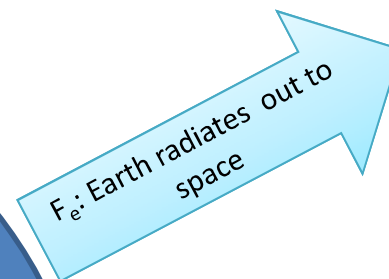
What is radiative forcing in a nutshell?



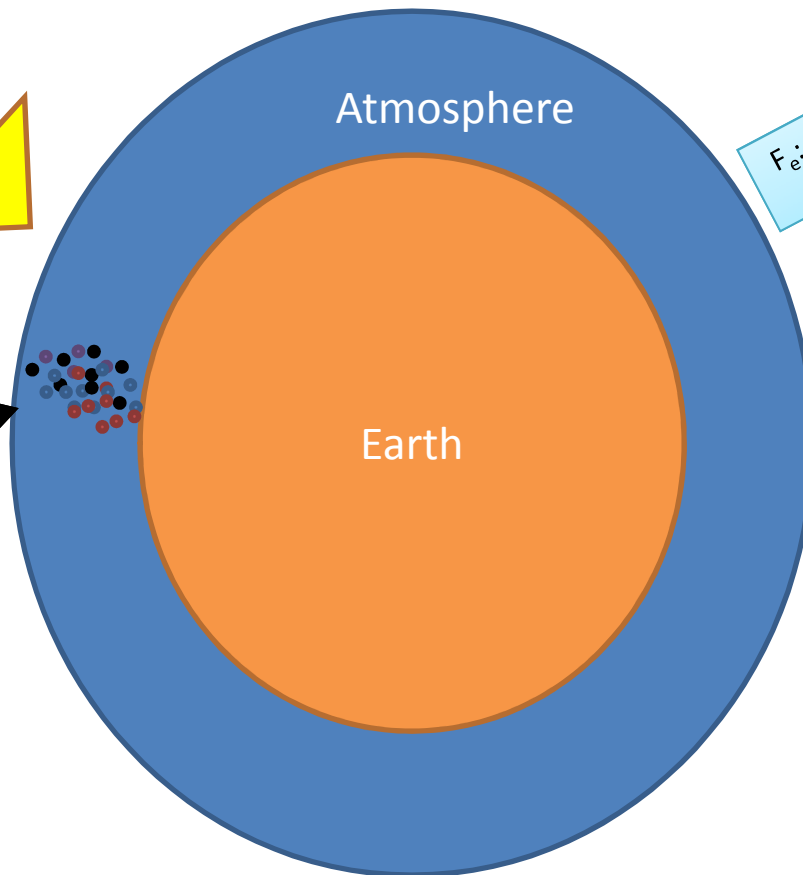
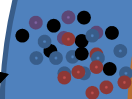


In equilibrium:

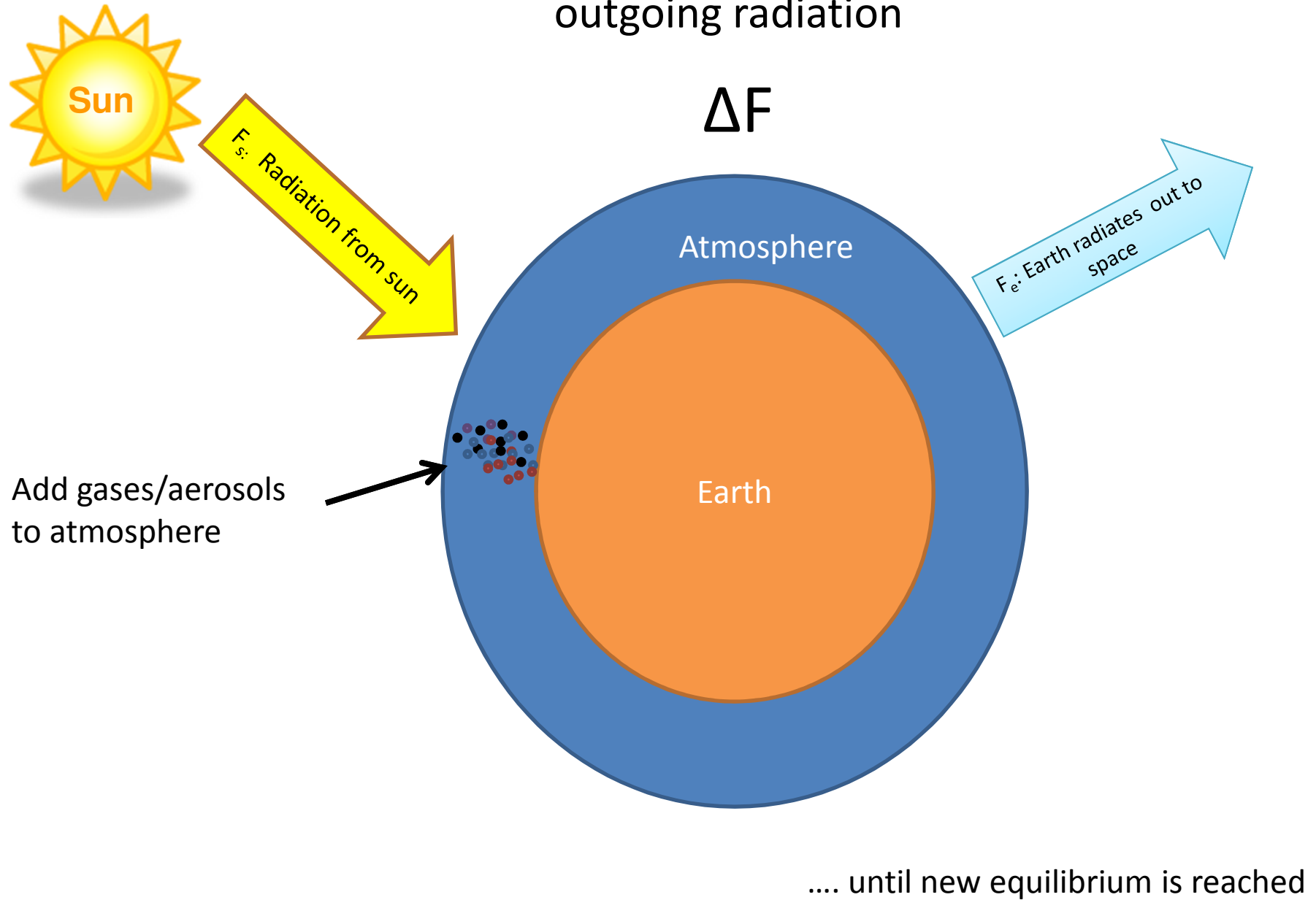
$$F_s = F_e$$



Add gases/aerosols  
to atmosphere

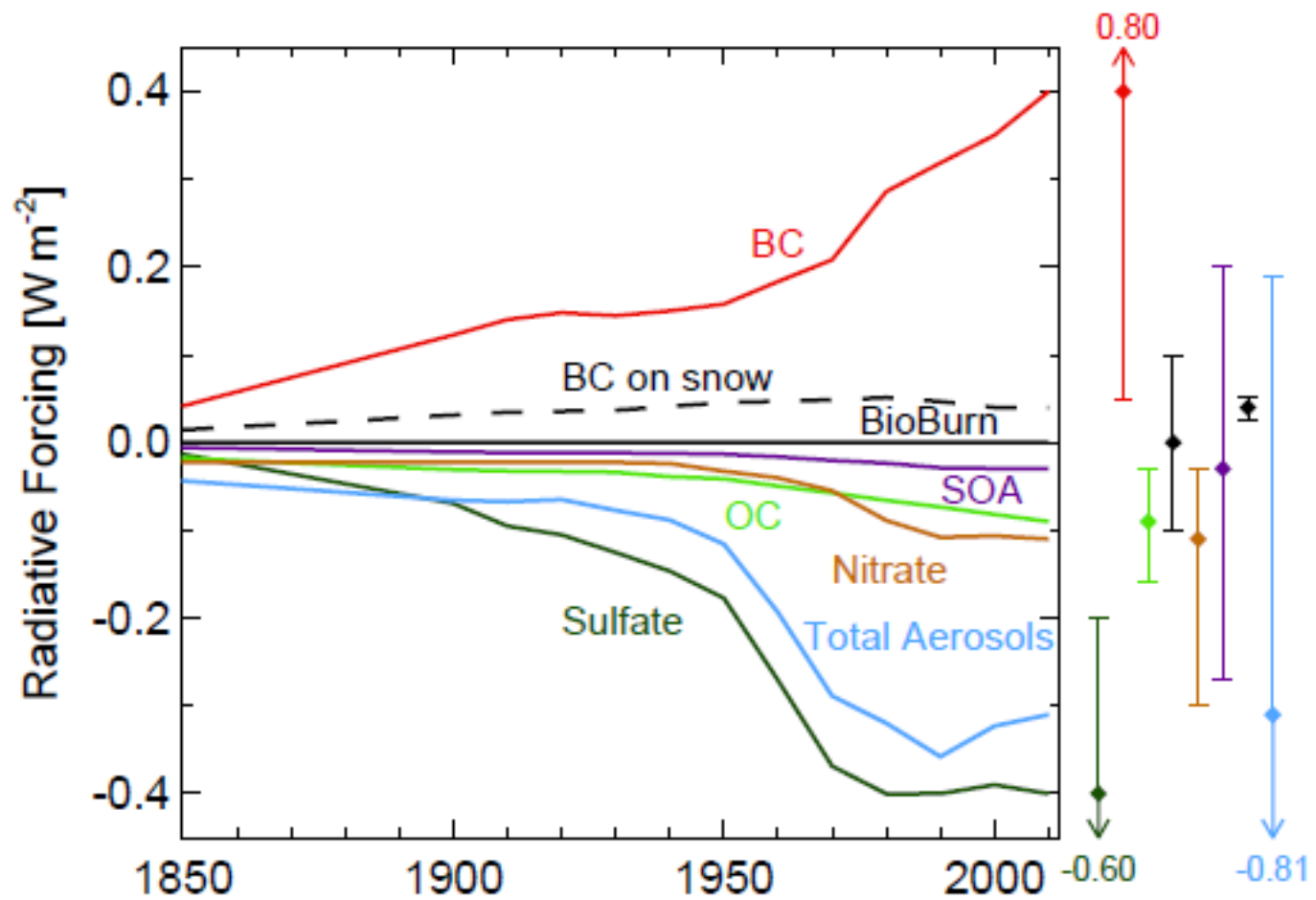


# Imbalance between incoming and outgoing radiation

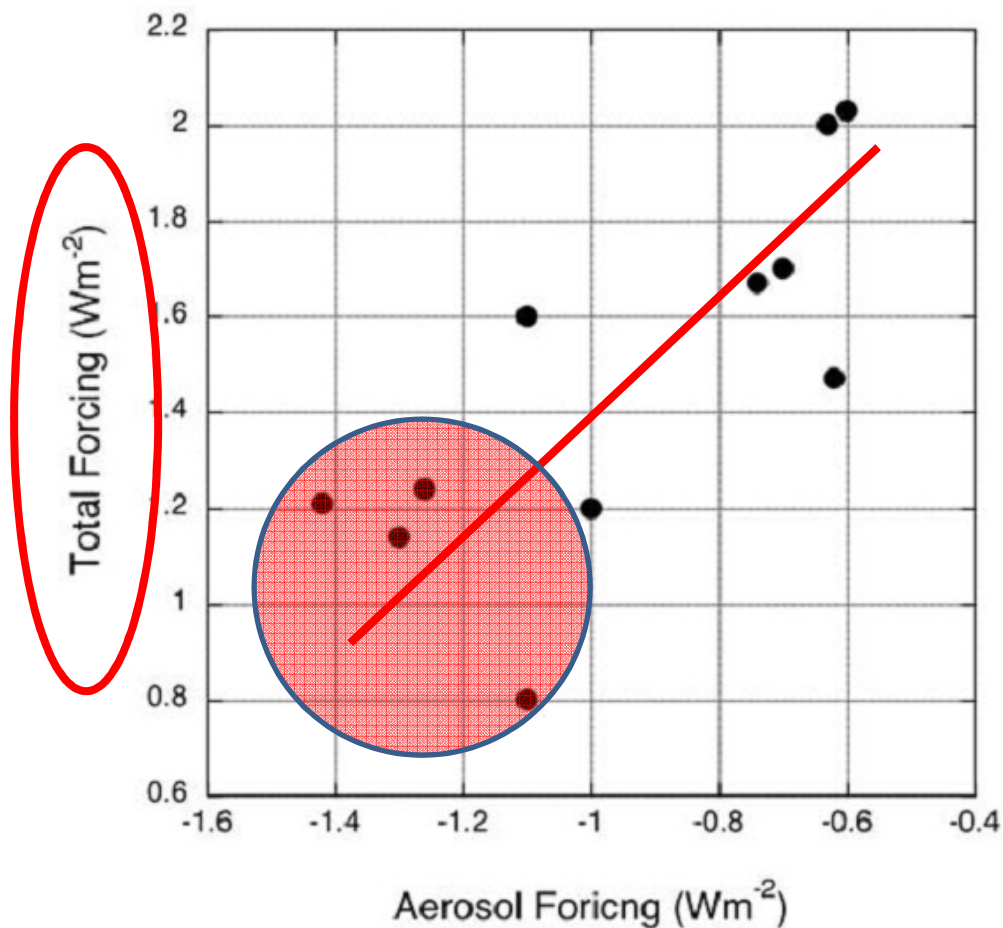




# Radiative forcing of aerosols, IPCC, 5AR, 2013



# Aerosols and Climate Sensitivity



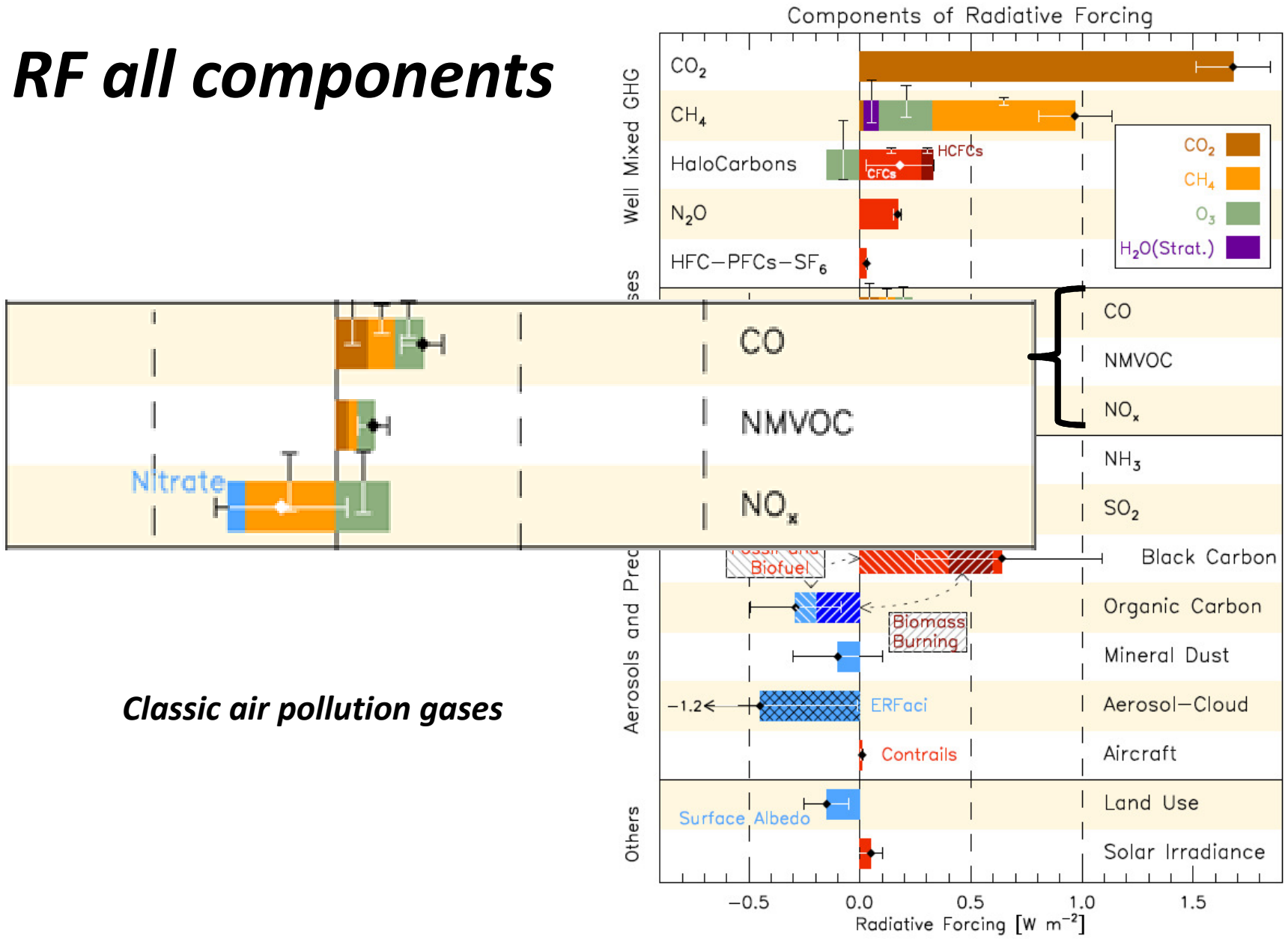
Climate Models are tuned by adjusting aerosol forcing, a poorly constrained process

Models with large climate sensitivity need large aerosol forcing

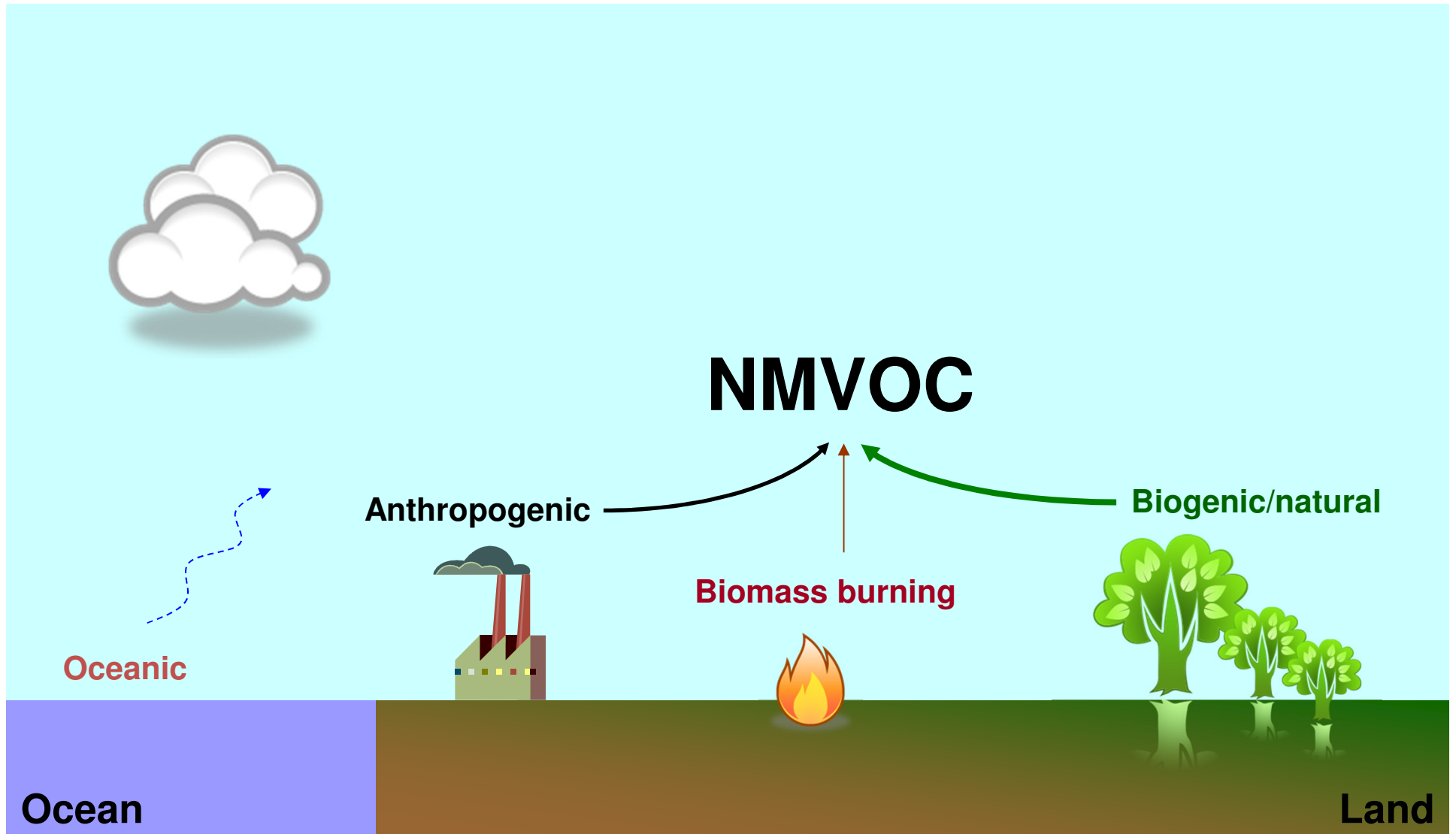
**Figure 2.** Total anthropogenic forcing ( $\text{Wm}^{-2}$ ) versus aerosol forcing ( $\text{Wm}^{-2}$ ) from nine fully coupled climate models and two energy balance models used to simulate the 20th century.

Kiehl, GRL, 2007

# RF all components



# Emission of NMVOC



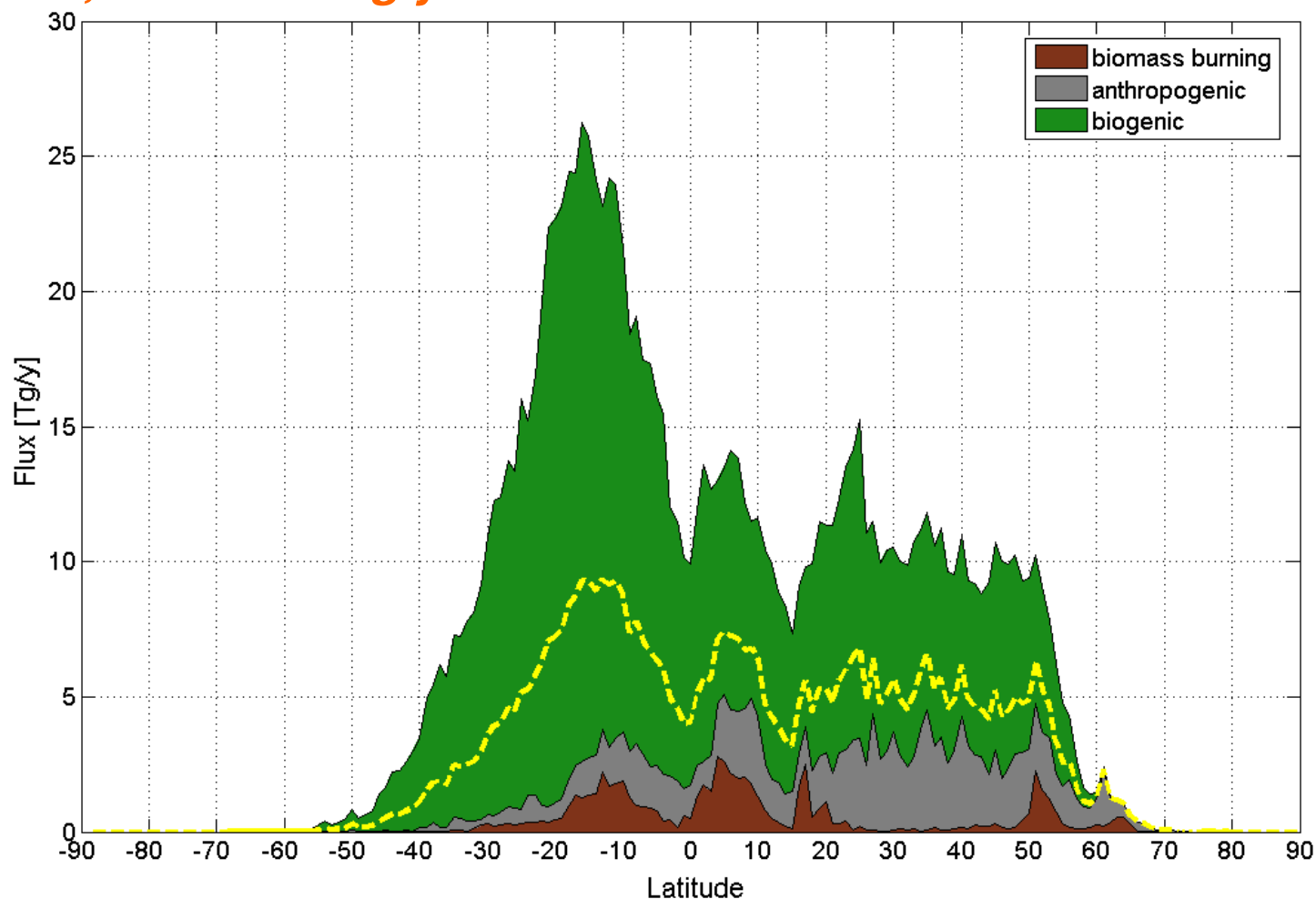
# Latitudinal Distribution of NMVOC

**Bottom-up (total): 1580 Tg/y**

**IPCC2001, 2007: 450 Tg/y**

Compare to CH<sub>4</sub>: about 500-600 Tg/y

Isoprene: about 500-700 Tg/y



A multi-investigator field mission in the SE USA in 2013

N<sub>itrogen</sub> O<sub>xidants</sub> M<sub>ercury</sub> A<sub>erosol</sub> D<sub>istribution</sub> S<sub>ources</sub> and S<sub>inks</sub>

S<sub>outh</sub> E<sub>ast</sub> NEX<sub>us</sub>

S<sub>outhern</sub> O<sub>xidants</sub> and A<sub>aerosol</sub> S<sub>tudy</sub>

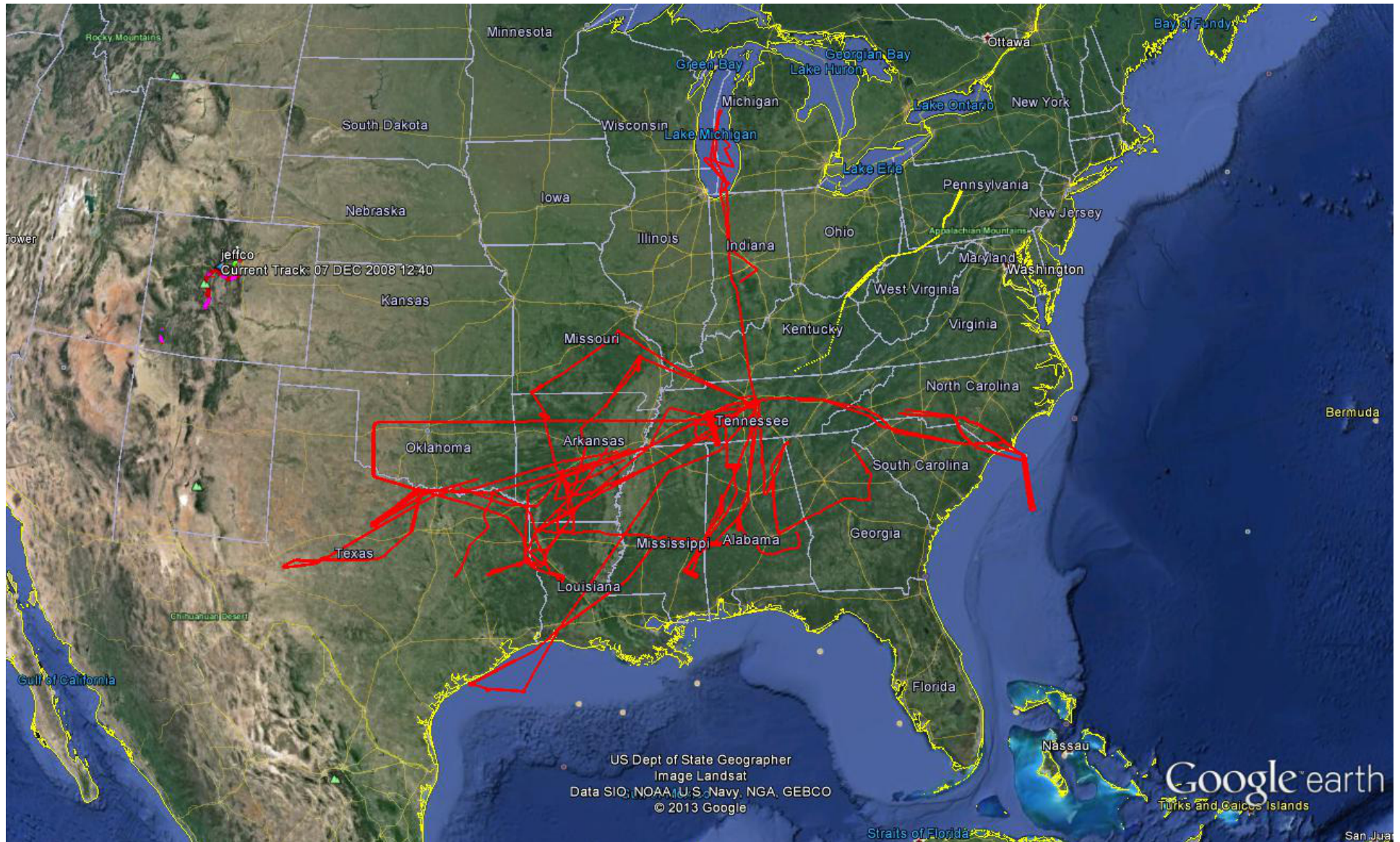
funded by NSF, NOAA, EPA

<http://catalog.eol.ucar.edu/sas>

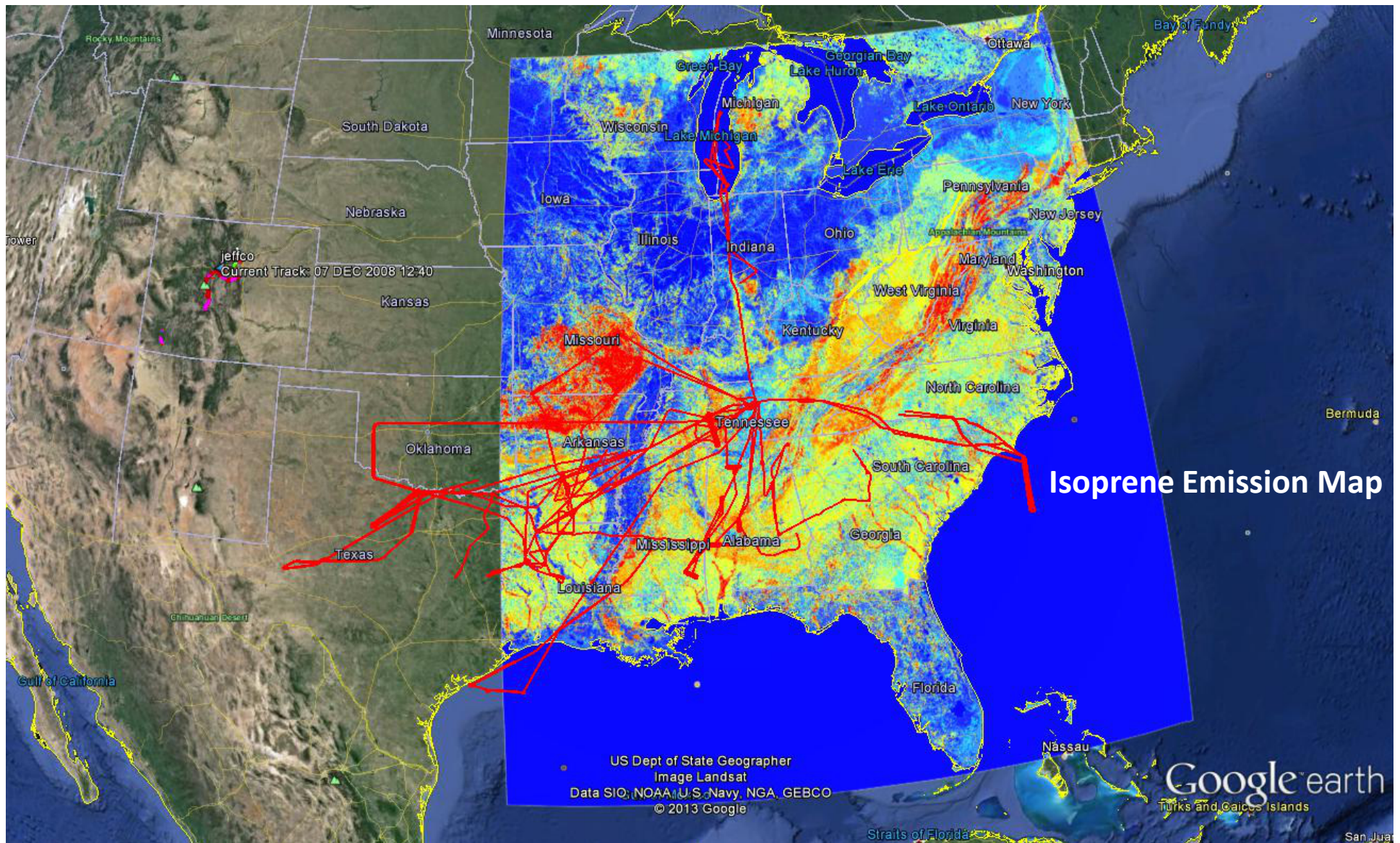
2013



# NOMADSS (Nitrogen, Oxidants, Mercury and Aerosol Distributions, Sources and Sinks SAS (Southeast Atmosphere Study) – *All flights*

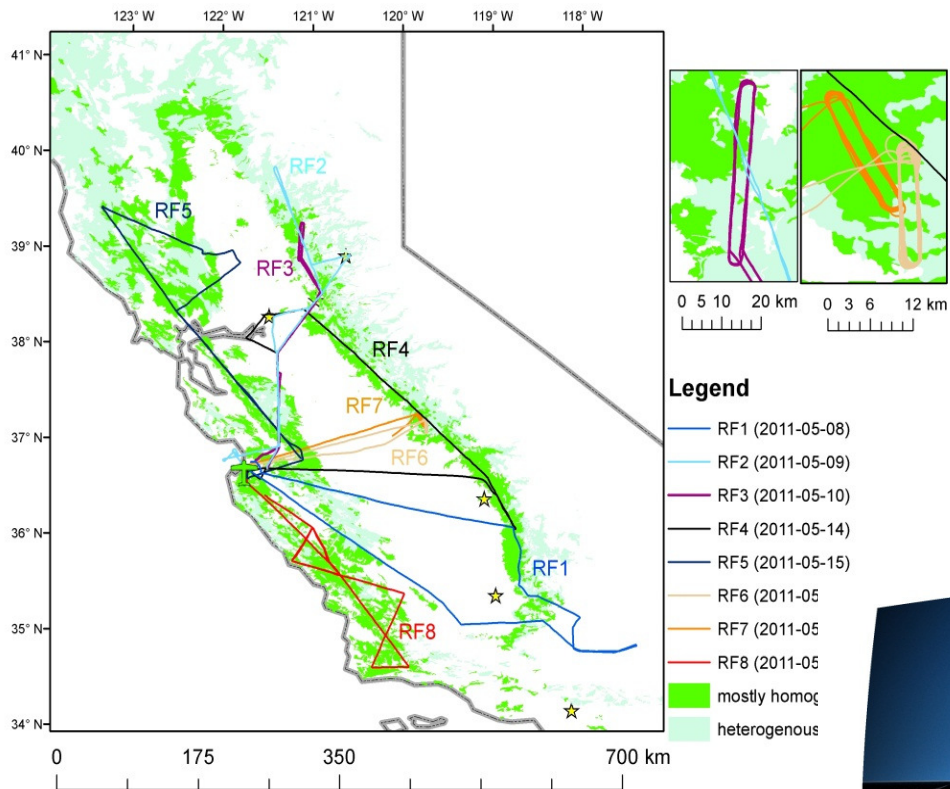


NOMADSS (Nitrogen, Oxidants, Mercury and Aerosol Distributions, Sources and Sinks  
SAS (Southeast Atmosphere Study) – **MEGAN Emission Factor MAP**





# California Airborne NMVOC Emission Research in Natural Ecosystem Transects (CABERNET) experiment, 2011



## Investigating Isoprene Fluxes

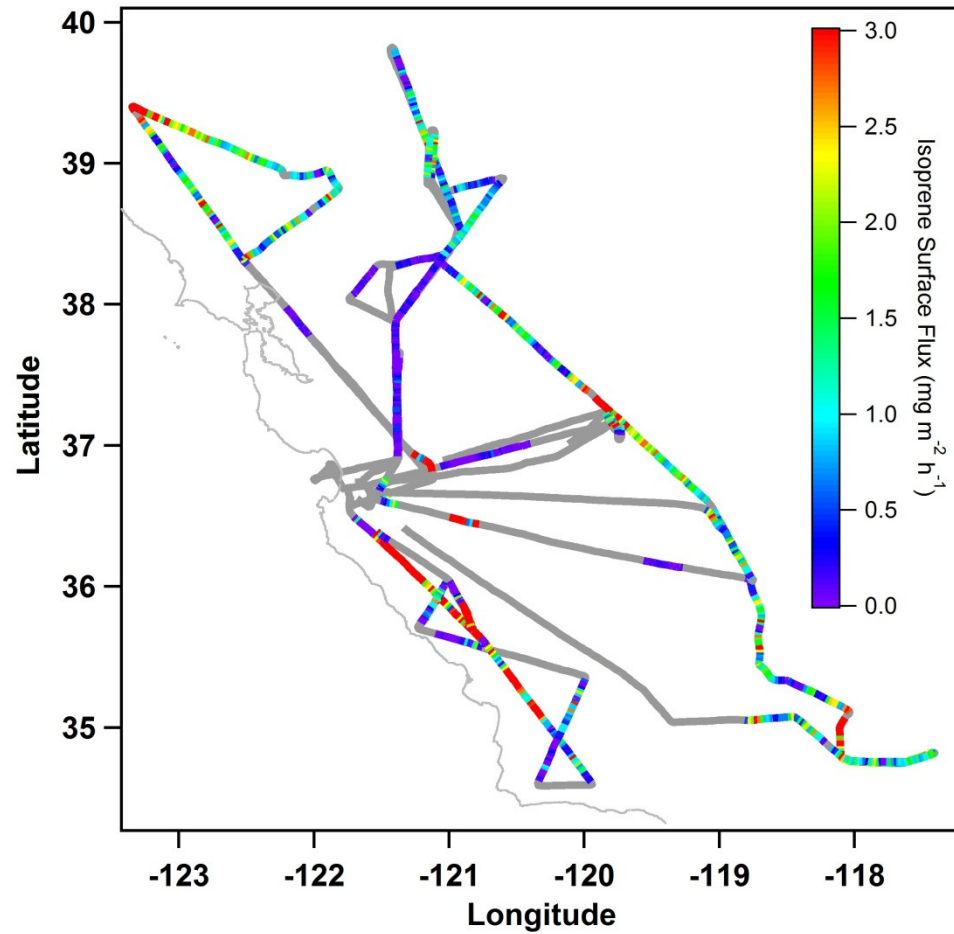


Karl et al., JAS, 2013

# Observationally derived spatially segregated NMVOC emission maps

Goal:

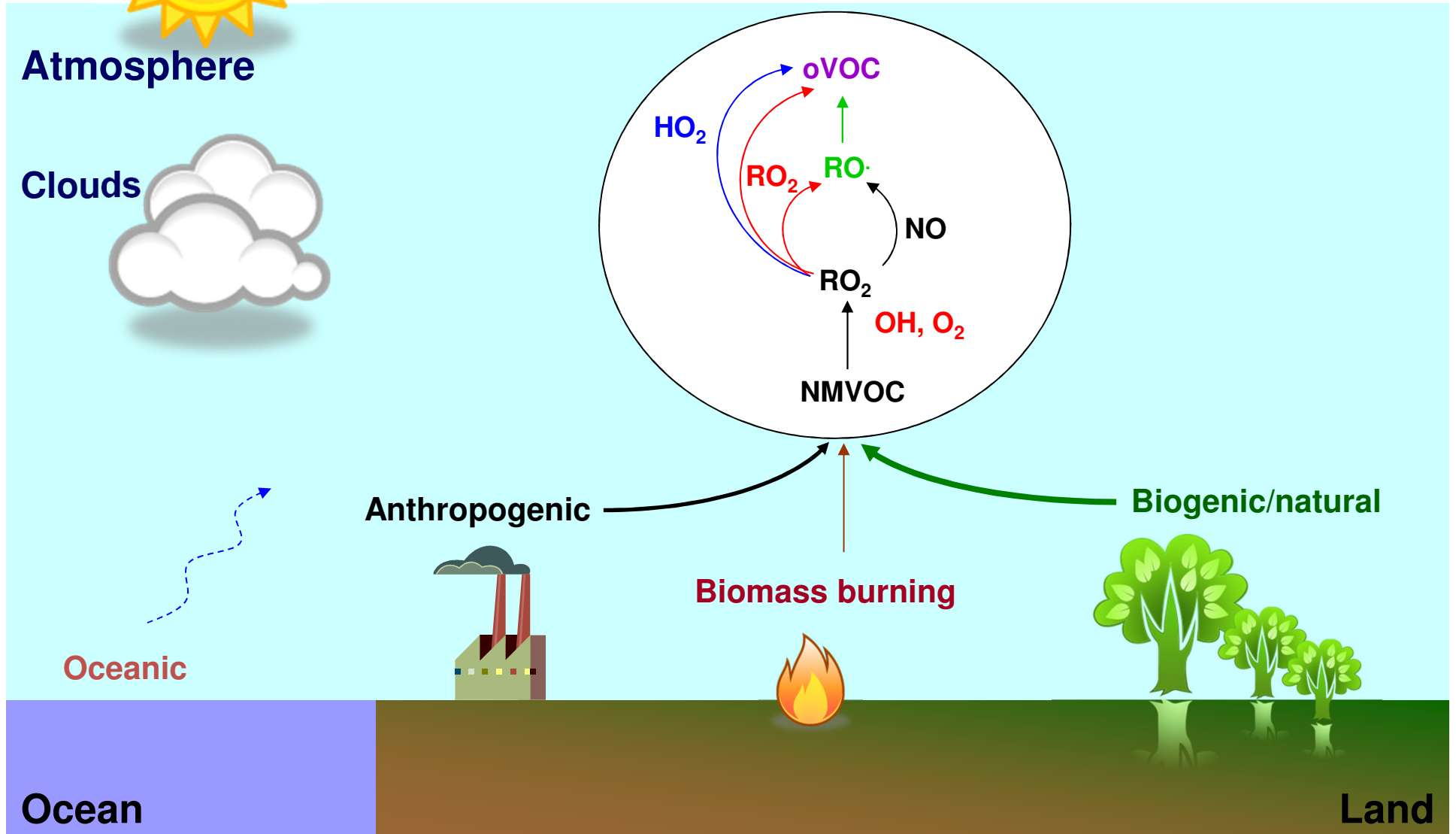
Dataset will be used to improve process based NMVOC models by constraining landscape level emission factors



# Earth System



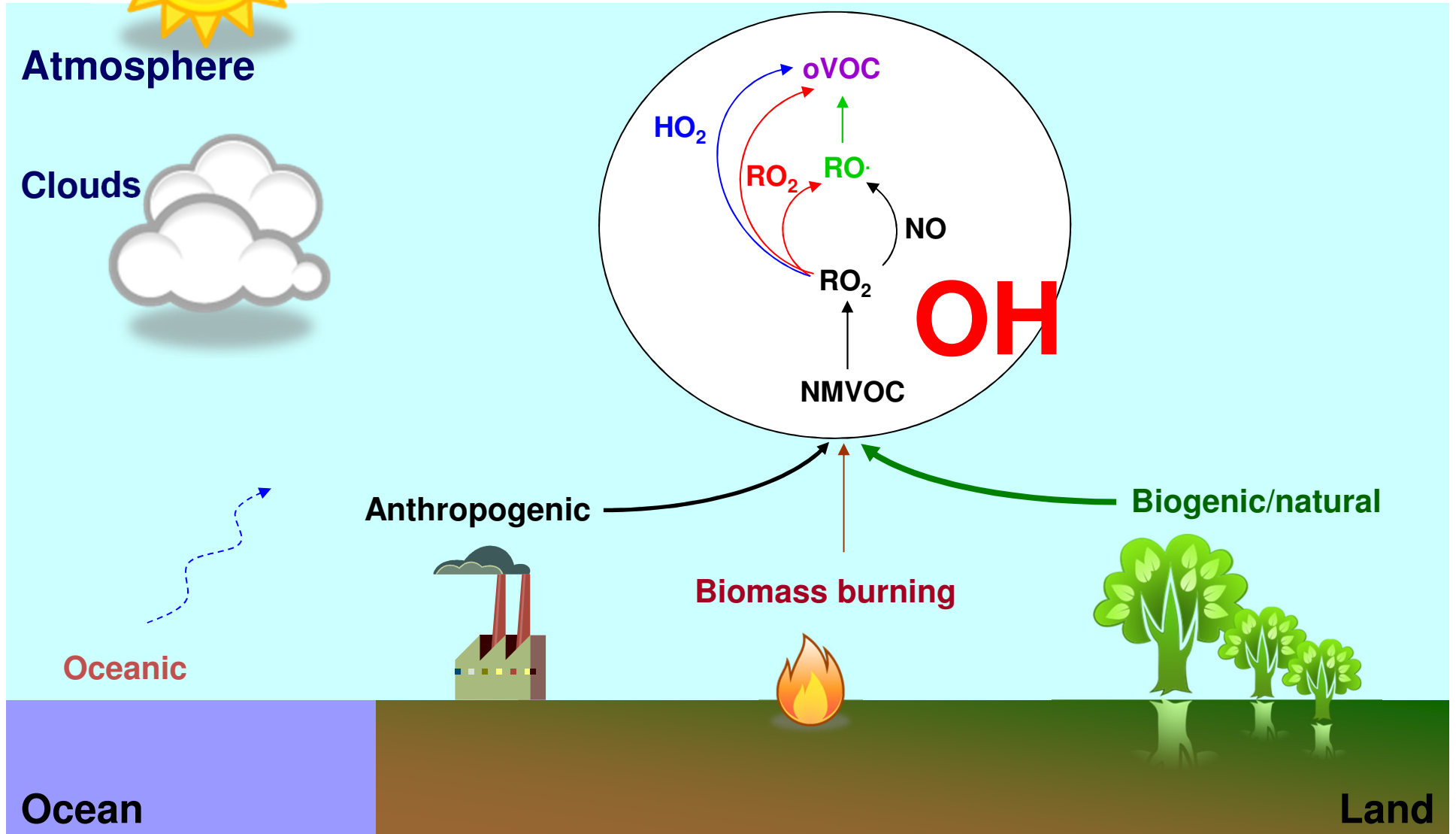
## Oxidation capacity of the atmosphere



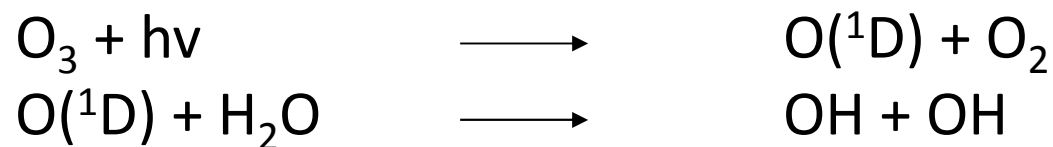
# Earth System



## Oxidation capacity of the atmosphere



# OH – The Detergent of the Atmosphere



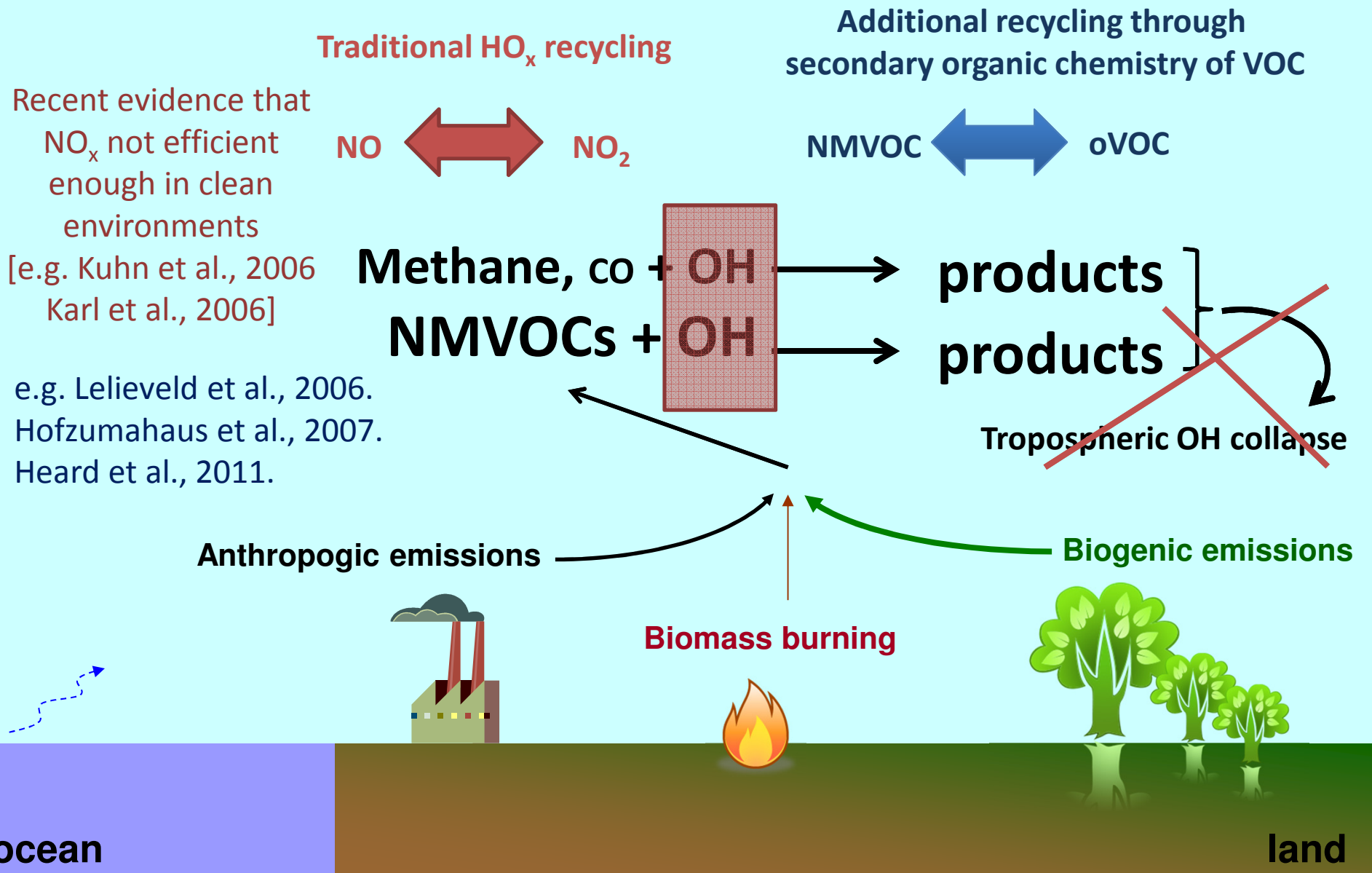
other HOx sources: photolysis of carbonyls, ozonolysis of hydrocarbons, OH+peroxides,  $\text{O}_3 + \text{HO}_2$

Very low mixing ratios:  $1 - 4 \times 10^{-14}$  (ppqv range)

Very reactive free radical – together with H and  $\text{HO}_2$  it forms the  $\text{HO}_x$  pool (= H + HO +  $\text{HO}_2$ )

UV radiation can not dissociate  $\text{O}_2$  for  $\text{HO}_x$  production in the troposphere anymore

# Chemical stability of the atmosphere



**Methane and isoprene (+monoterpenes) are the among the most important reactive carbon containing trace gases to understand paleoclimate**

## Enhanced chemistry-climate feedbacks in past greenhouse worlds

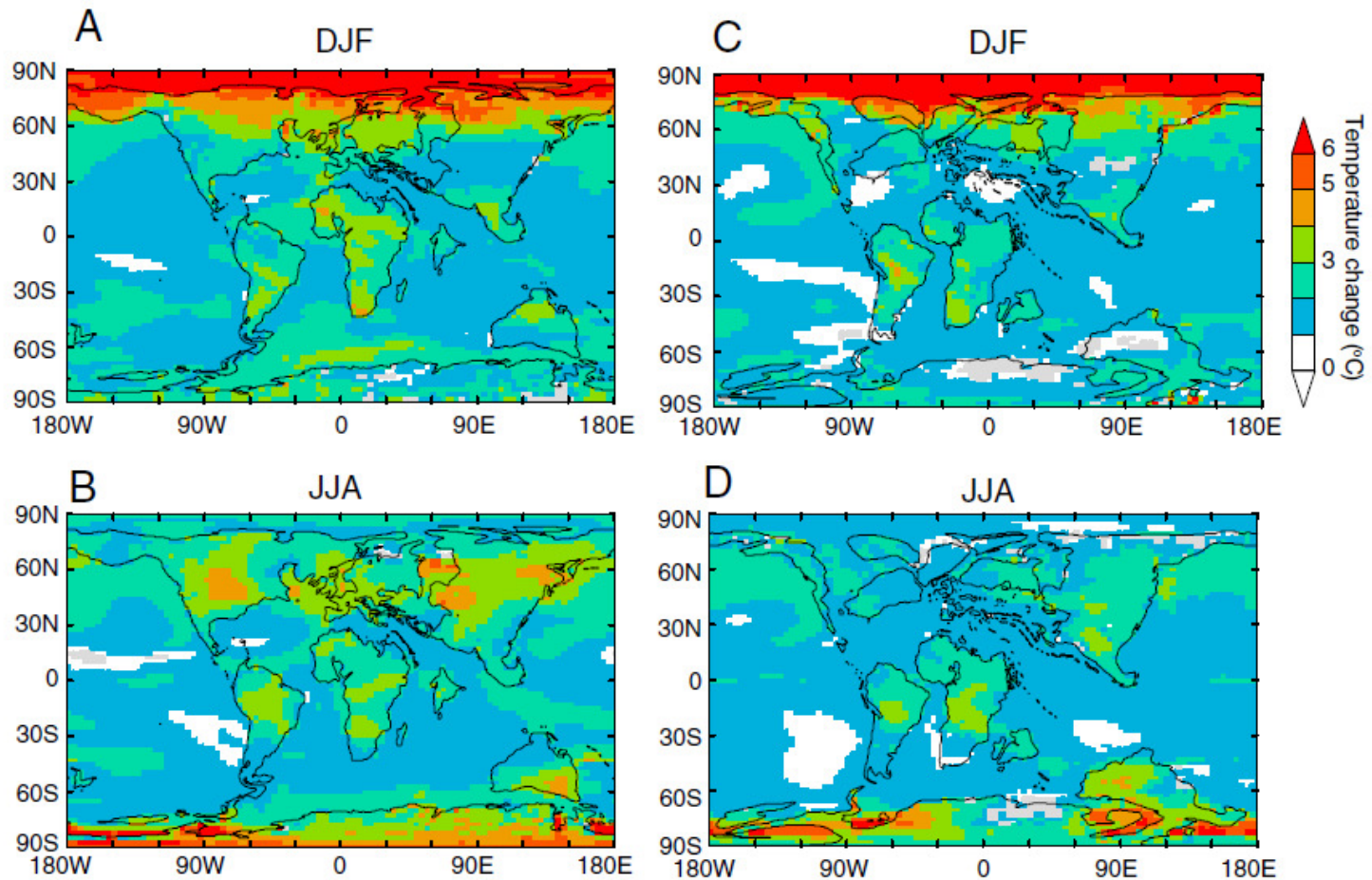


David J. Beerling<sup>a,1</sup>, Andrew Fox<sup>a,2</sup>, David S. Stevenson<sup>b</sup>, and Paul J. Valdes<sup>c</sup>

<sup>a</sup>Department of Animal and Plant Sciences, University of Sheffield, Sheffield S10 2TN, United Kingdom; <sup>b</sup>School of GeoSciences, University of Edinburgh, Edinburgh EH9 3JN, United Kingdom; and <sup>c</sup>Department of Geographical Sciences, University of Bristol, Bristol BS8 1SS, United Kingdom

Edited by Ralph J. Cicerone, National Academy of Sciences, Washington, DC, and approved April 26, 2011 (received for review February 11, 2011)

Isoprene controls the lifetime of methane – both emissions are much higher than pre industrial levels – this system causes a high climate sensitivity due to it's impact on OH



Eocene  
55 Ma

Cretaceous  
90 Ma

High climate sensitivity through chemistry – climate interactions

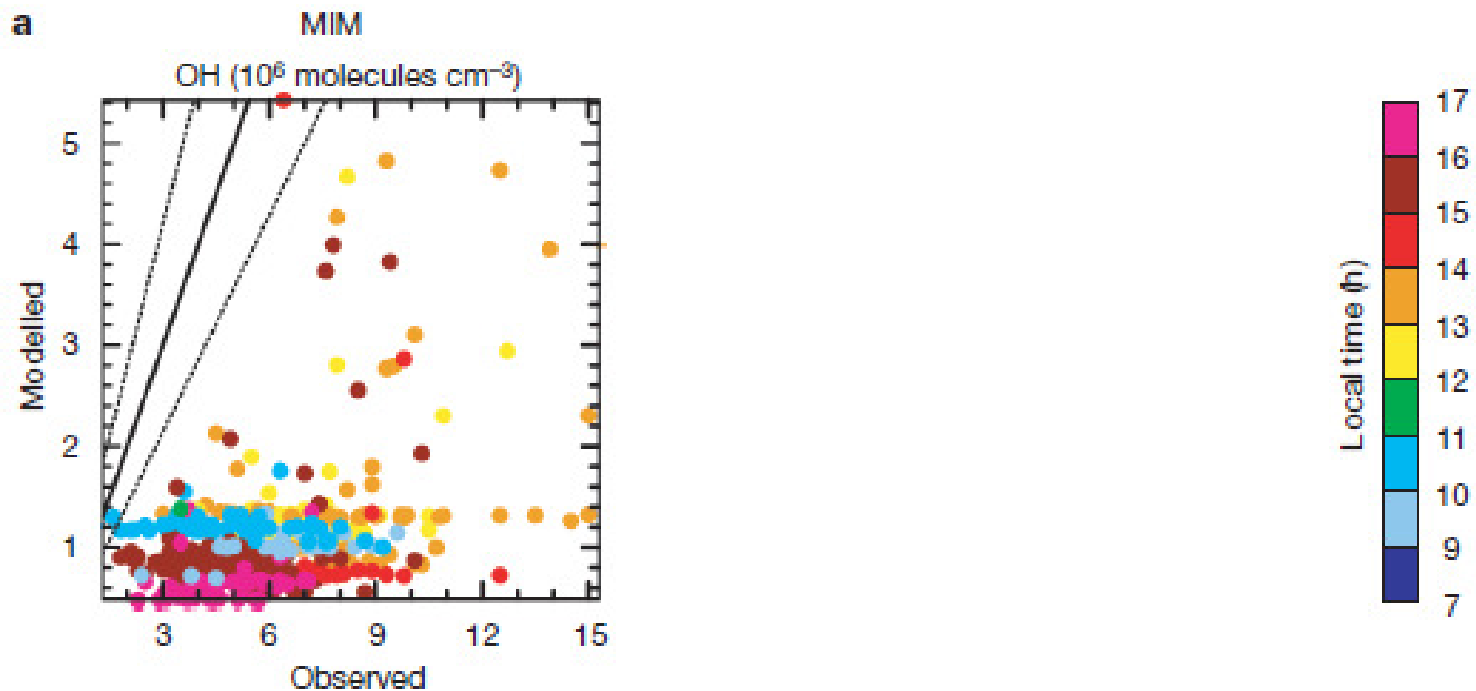


First assessment of the performance of a global CT Model

# Atmospheric oxidation capacity sustained by a tropical forest

Nature, 2008

J. Lelieveld<sup>1</sup>, T. M. Butler<sup>1</sup>, J. N. Crowley<sup>1</sup>, T. J. Dillon<sup>1</sup>, H. Fischer<sup>1</sup>, L. Ganzeveld<sup>1</sup>, H. Harder<sup>1</sup>, M. G. Lawrence<sup>1</sup>, M. Martinez<sup>1</sup>, D. Taraborrelli<sup>1</sup> & J. Williams<sup>1</sup>



Also: Hofzumahaus et al. , Science, 2009. Paulot et al. , Science, 2009.

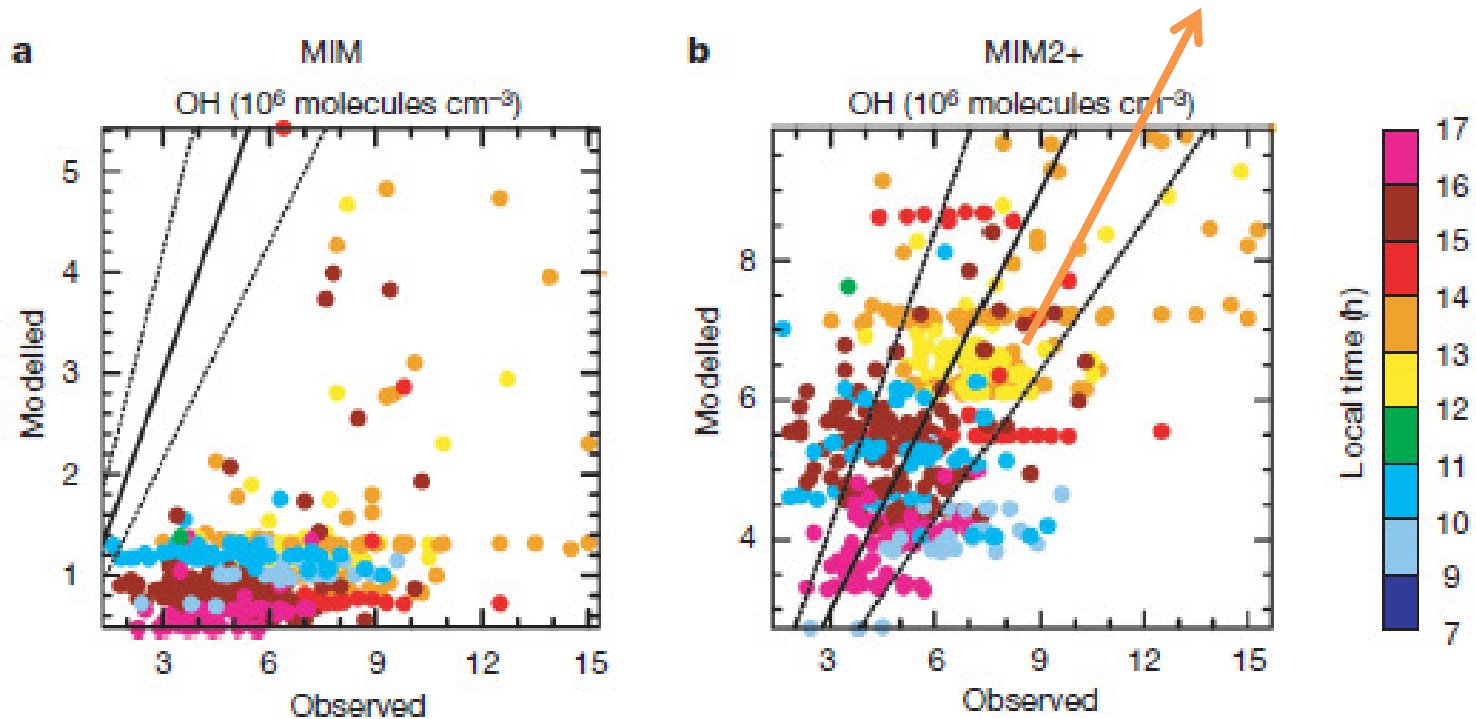
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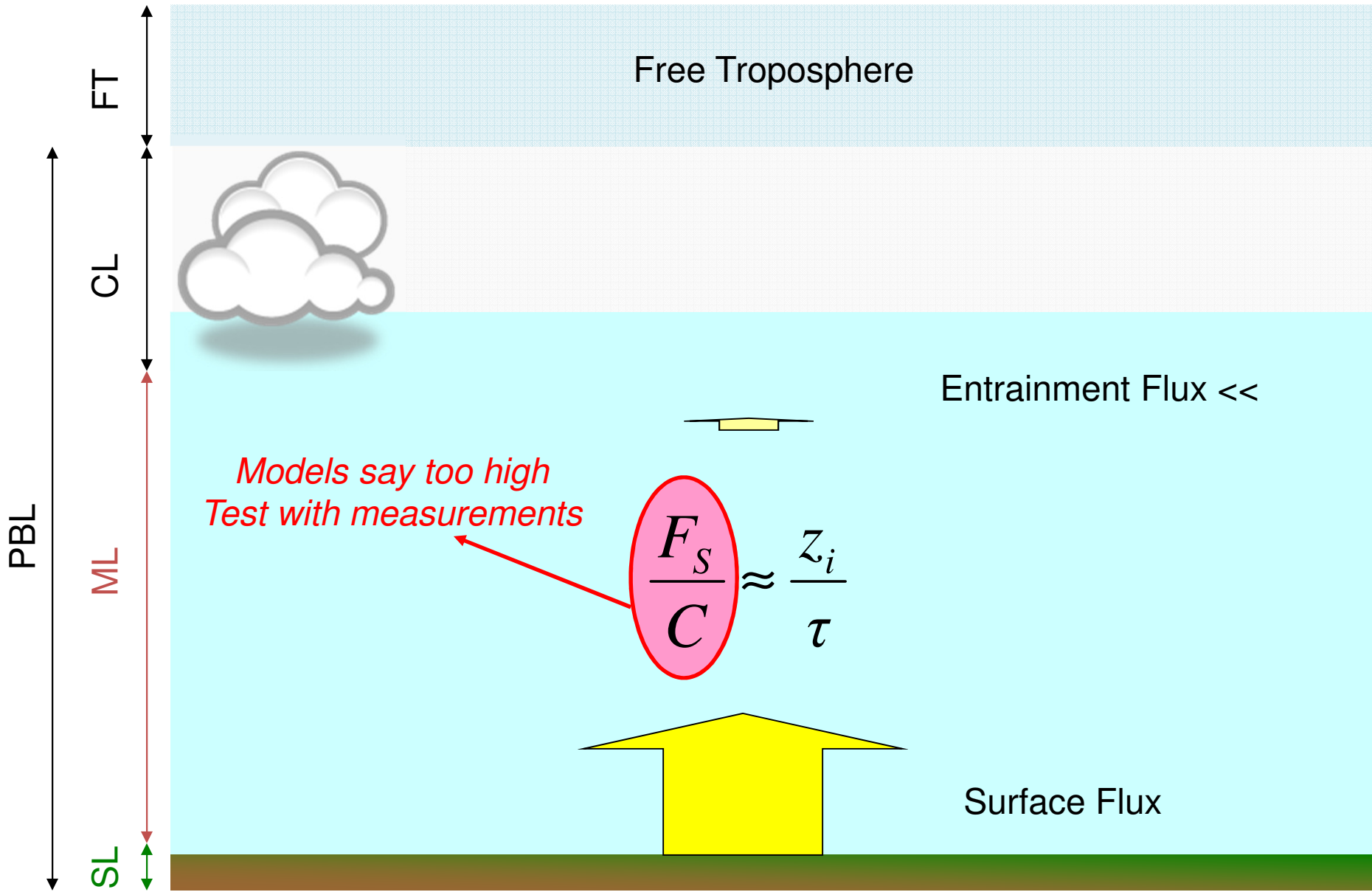
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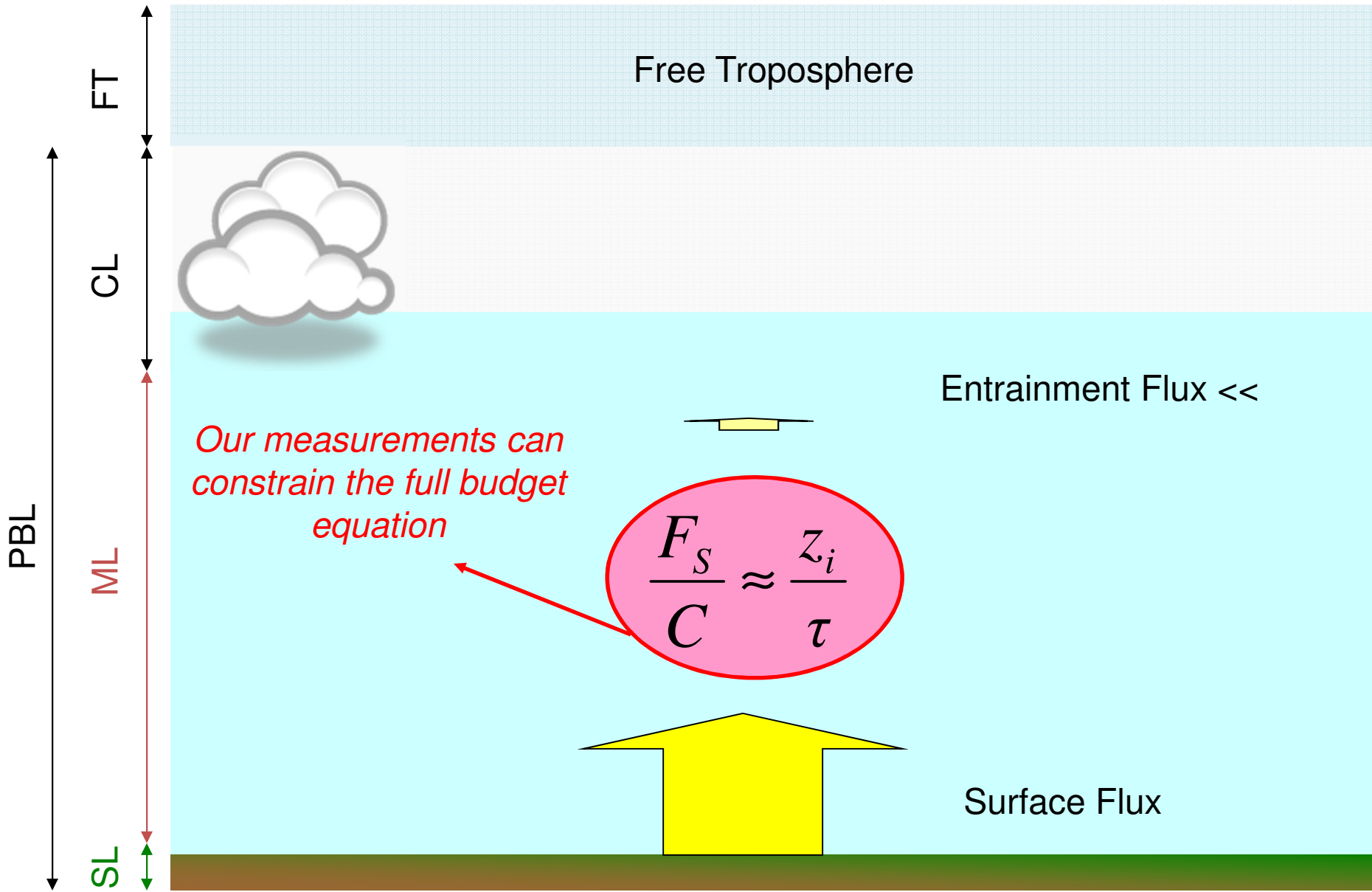
*Model chemistry „tuned“*



Also:

Hofzumahaus et al. , Science, 2009. Paulot et al. , Science, 2009. Taraborelli et al., Nat. Geo., 2012





Free Troposphere

FT

CL



Entrainment Flux <<

*Our measurements can constrain the full budget equation*

$$\frac{F_s}{C} \approx \frac{z_i}{\tau}$$

Surface Flux

PBL

ML

SL

# CABERNET - results

Minimum of sensible heat flux

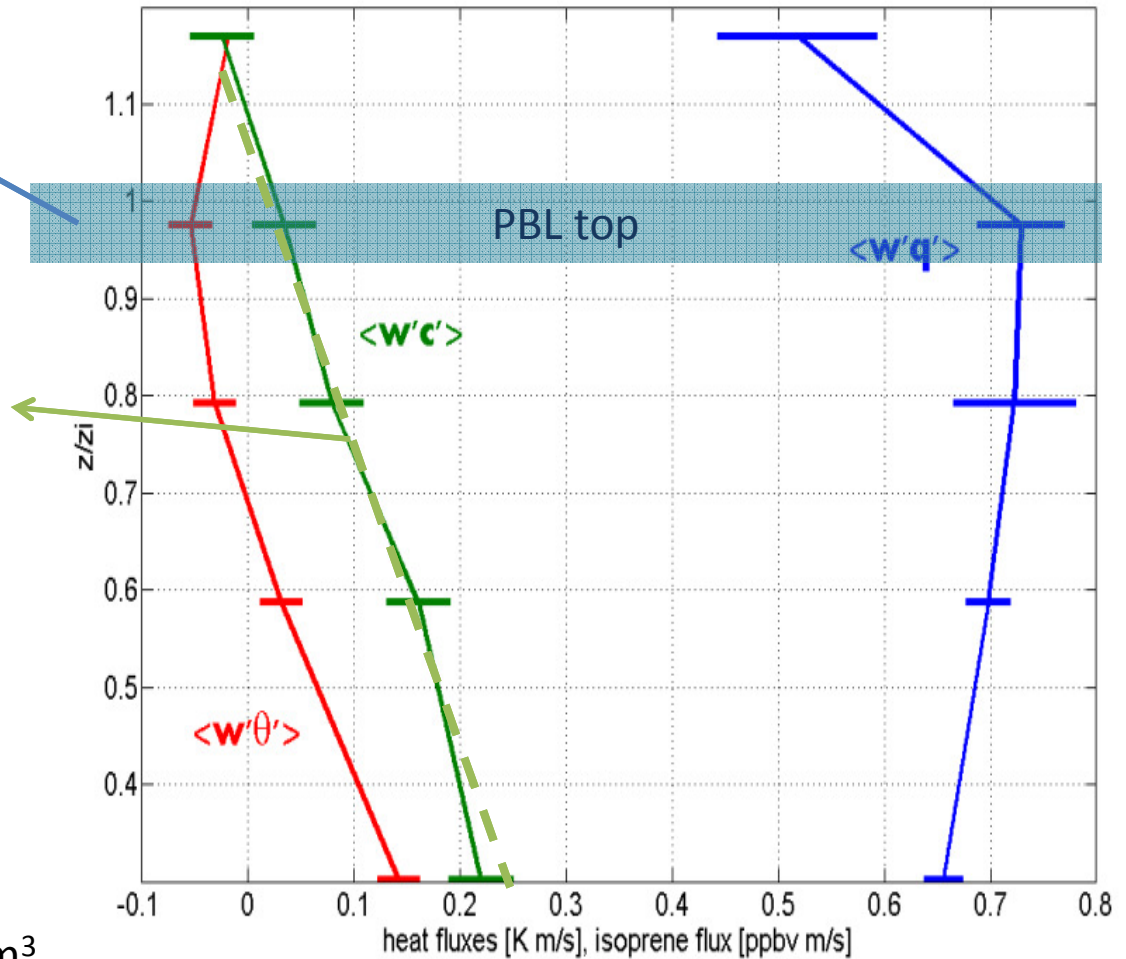
Slope of Isoprene flux divergence is proportional to Damkohler (Da) number and the *lifetime* imposed by the *OH radical*

3 research flights:

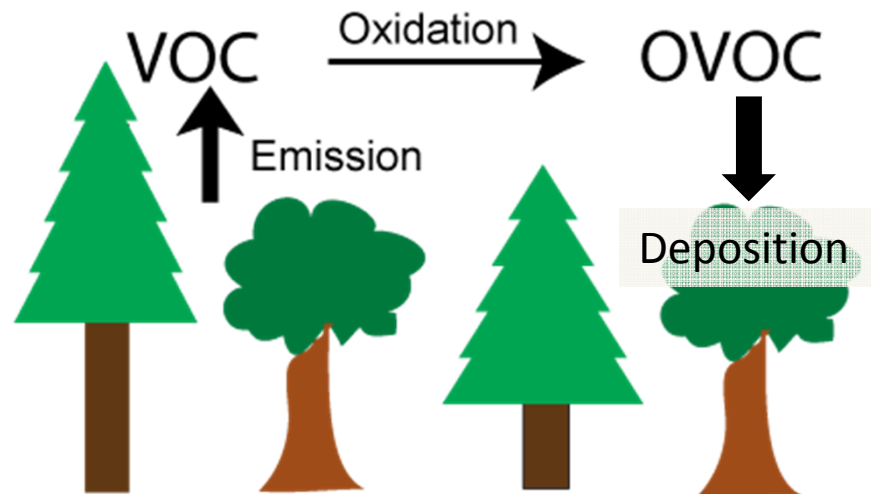
Da: 0.2 – 0.9

$v_e$  : 1.5 – 9.6 cm/s

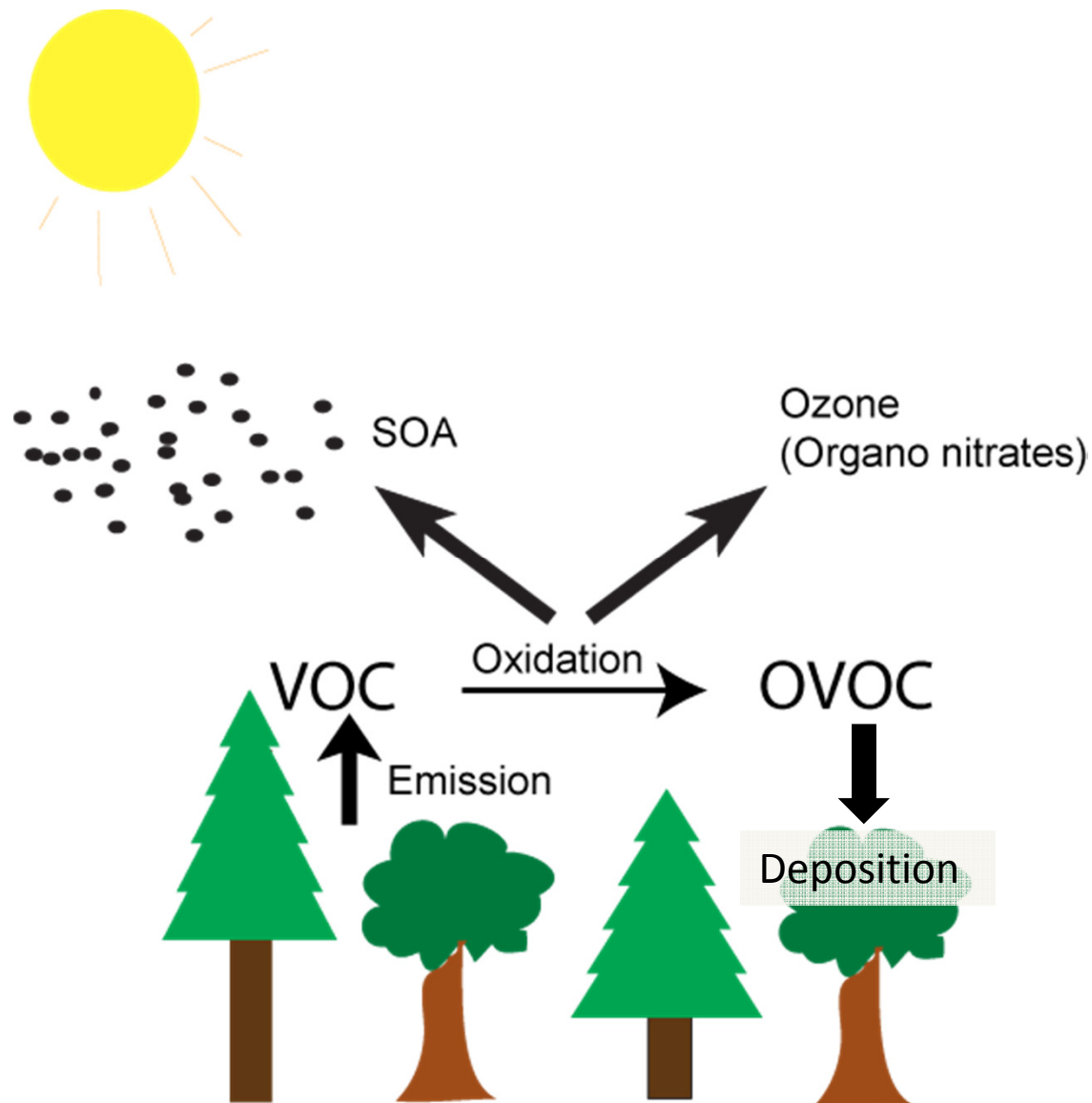
OH: 4.4 – 7.2 x 10<sup>6</sup> molecules/cm<sup>3</sup>



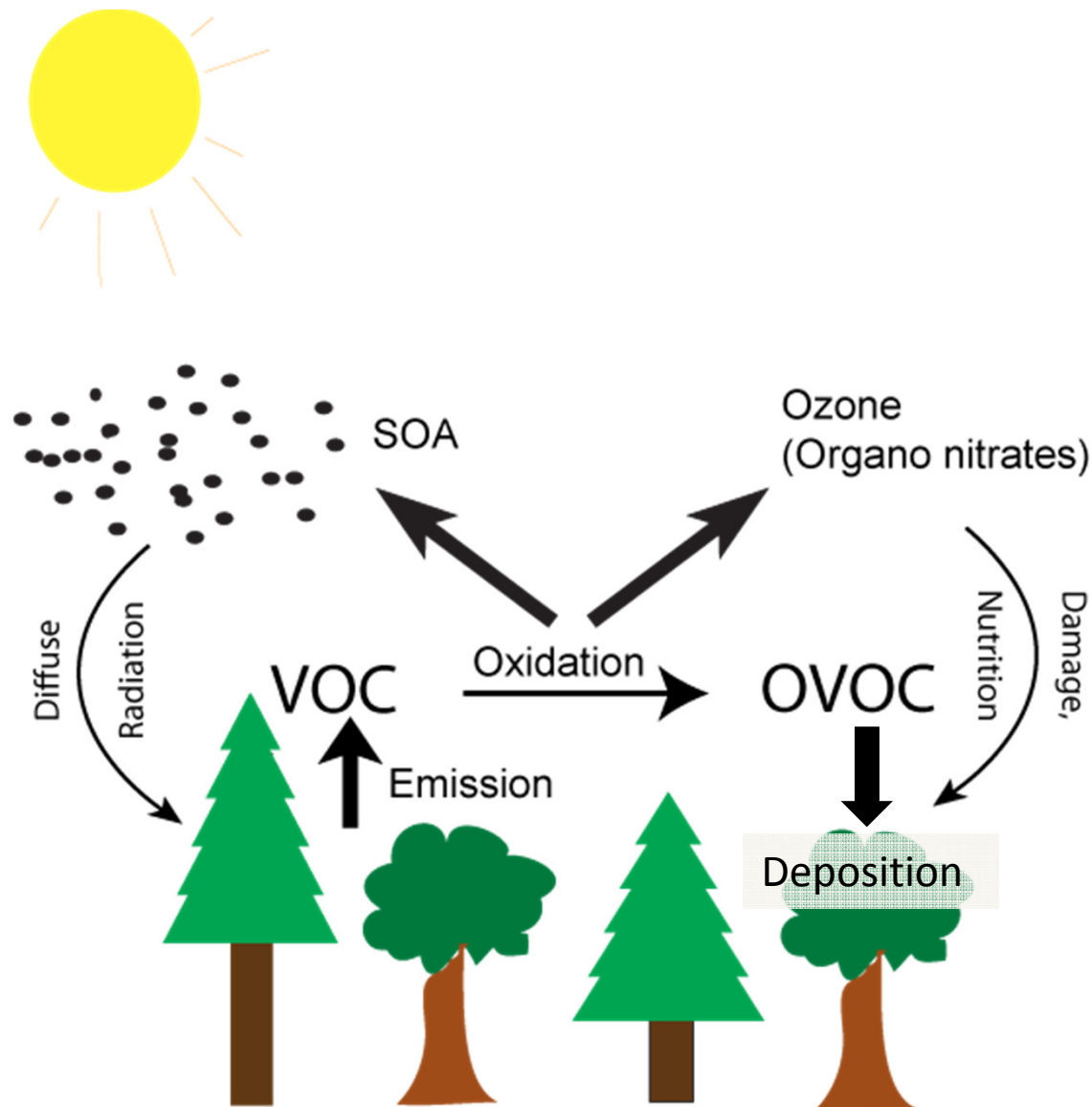
# NMVOC Oxidation, Ozone and SOA Formation: Air pollution-Atmosphere Feedbacks



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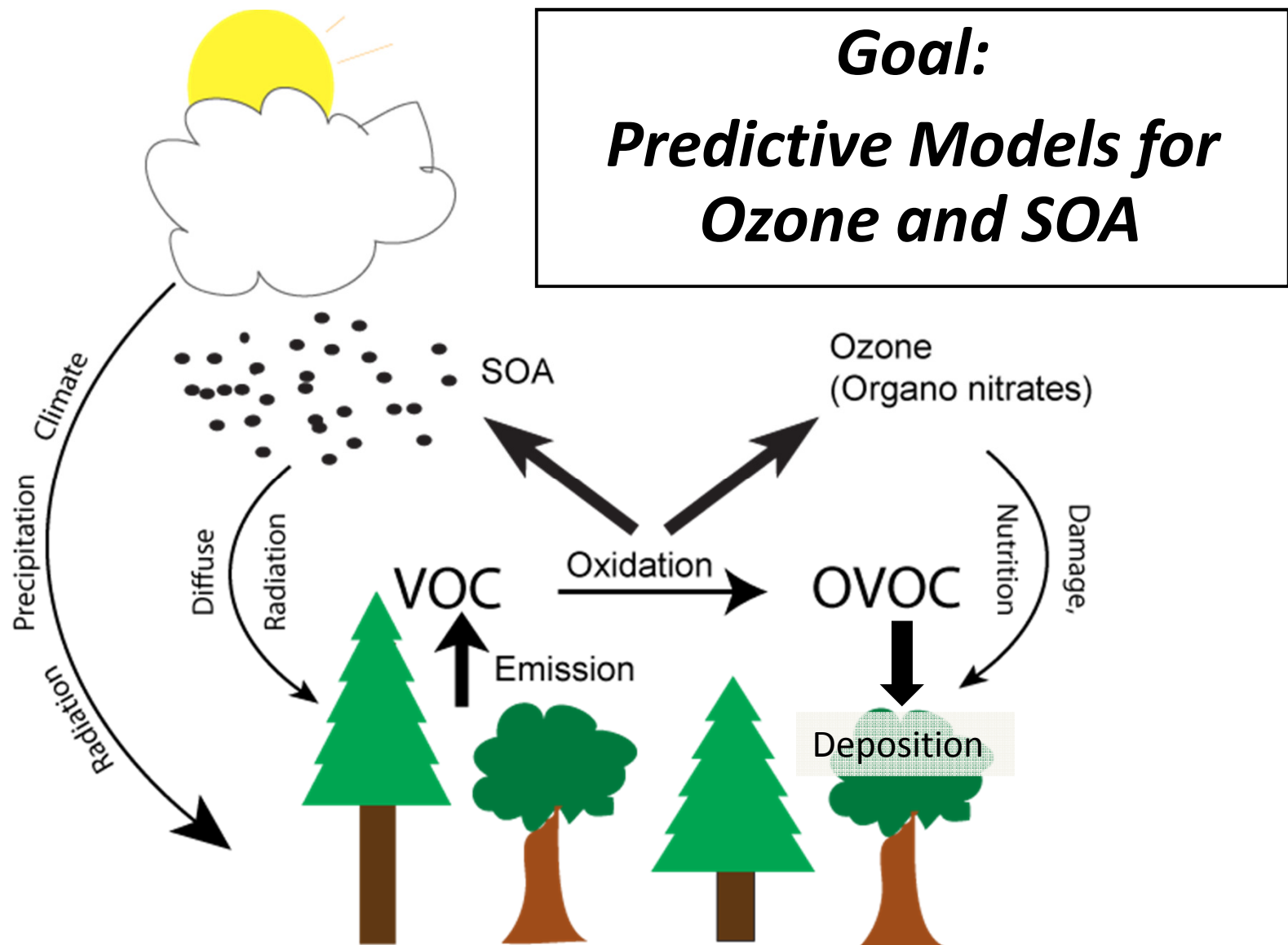


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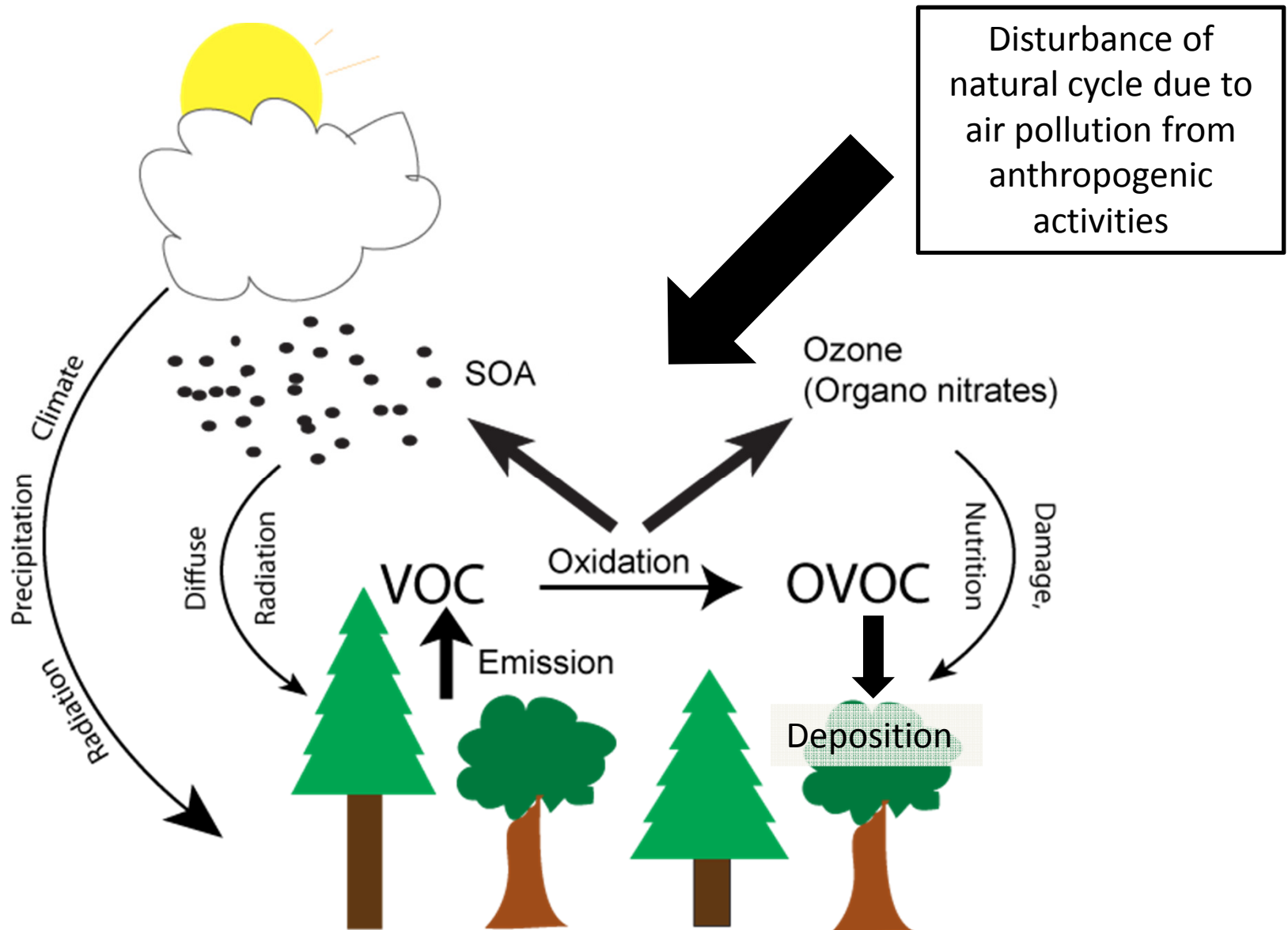




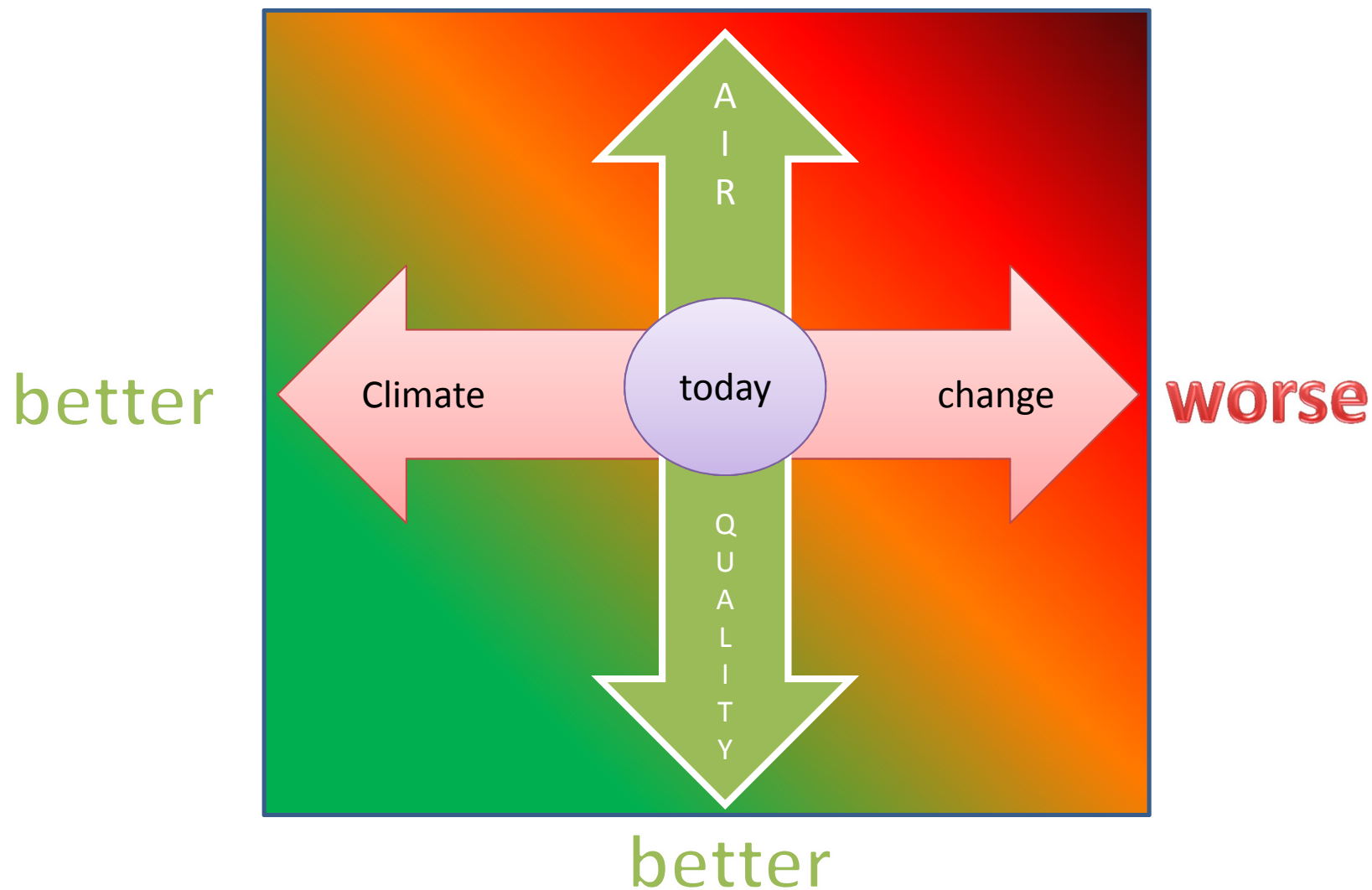
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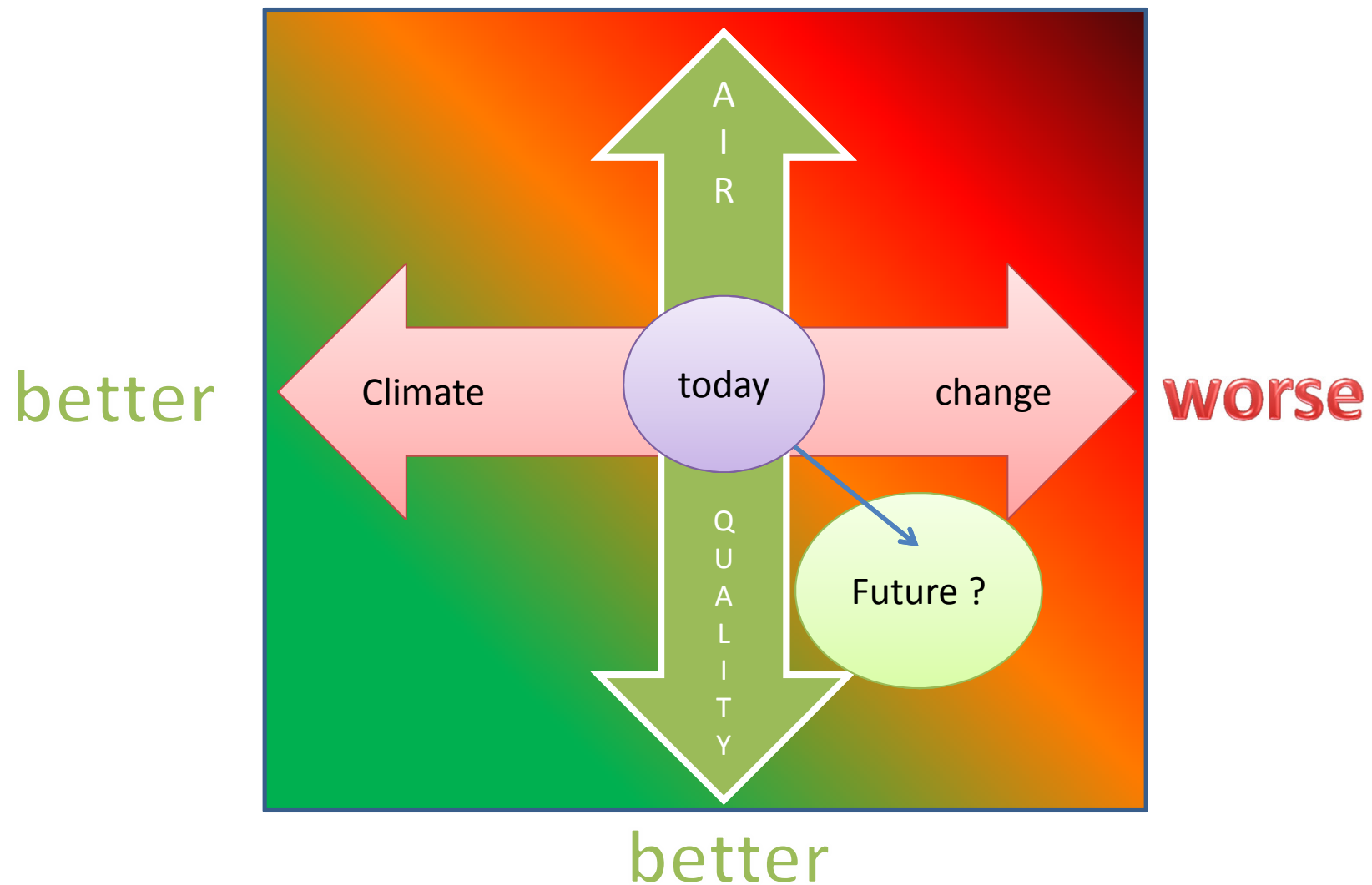
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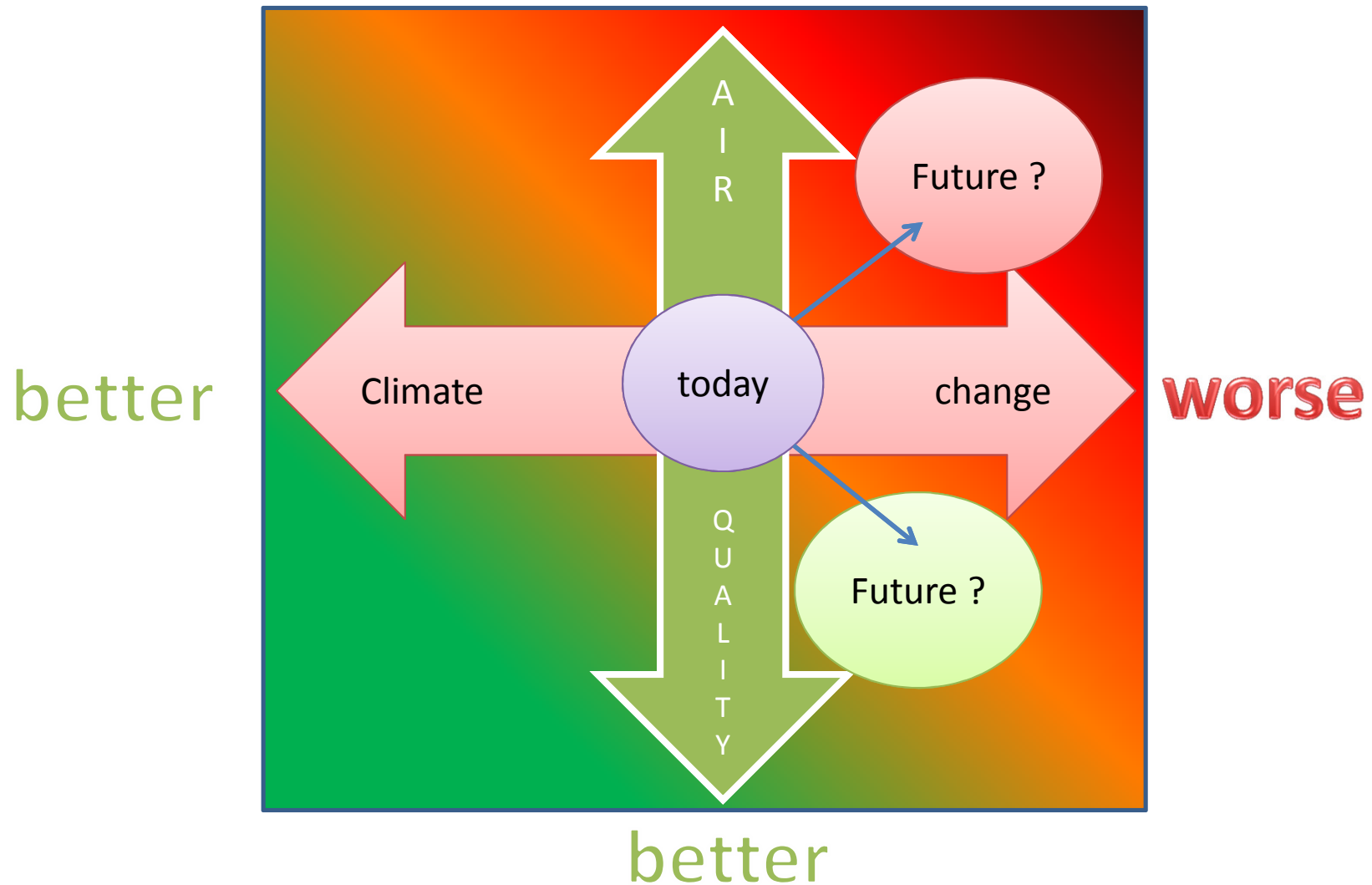
***Science to support decisions: Climate vs Air Quality  
(AQ) – at a cross roads  
worse***

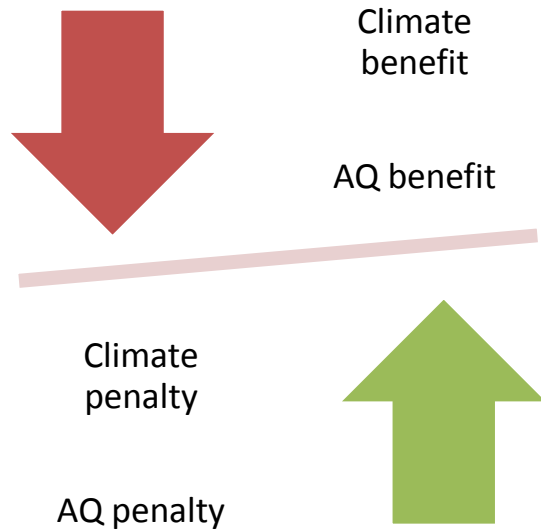


***Science to support decisions: Climate vs Air Quality  
(AQ) – at a cross roads  
worse***

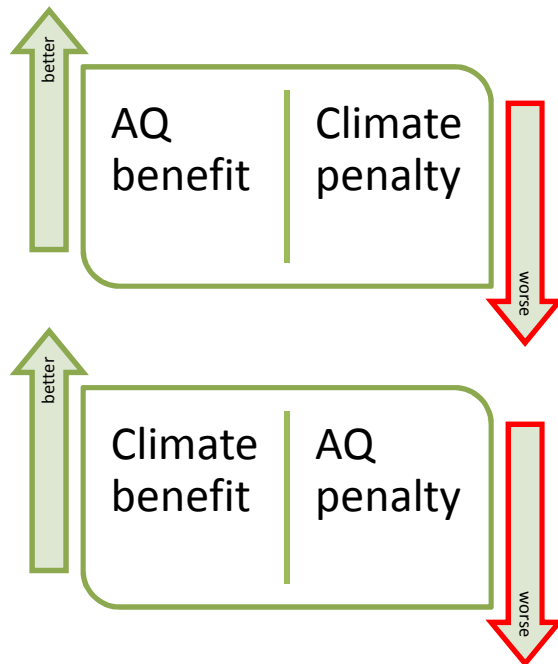


*Science to support decisions: Climate vs Air Quality  
(AQ) – at a cross roads  
worse*





Win-win: AQ and climate mitigation strategies benefit from each other: an ideal scenario for policy makers



However relationships between AQ and climate can also behave antagonistically

# AQ-climate interaction

Climate mitigation exhibits a complex interaction with air quality control measures

As an example the National Academy of Sciences (NAS) recommended in its 2004 report, *Air Quality Management in the United States*, that air pollution and climate change policies be developed through an integrated approach.

# Climate penalty on Air Quality

- ***Rising temperature*** will result in enhanced emissions of biogenic ***NMVOCs*** (volatile organic compounds) -> increase in ozone and SOA
- Longer stagnant periods/heat waves will ***accumulate pollutants***
- Biomass fuels: increase in black carbon and SOA – could also increase in regional NMVOC
- .....



# Air Quality penalty on Climate

- Reduction of aerosols
  - Reduction of primary aerosols (other than black carbon)
  - Reduction of anthropogenically enhanced biogenic SOA formation
  - Reduction of SO<sub>2</sub>

# Summary

- Reducing some short-lived constituents have a strong immediate climate benefit
- AQ-climate interactions can be complex and need to be considered for policy making
- Feedbacks can dampen or enhance response processes in the climate system (e.g. higher temperatures -> higher natural NMVOC emissions - > more SOA -> negative forcing - more cooling?)
- Overall it is expected that AQ will increasingly suffer due to climate change (e.g. longer stagnant periods such as the 2003 heat wave in Europe)
  - This might mandate stricter AQ measures to maintain the current status (Climate penalty on AQ)



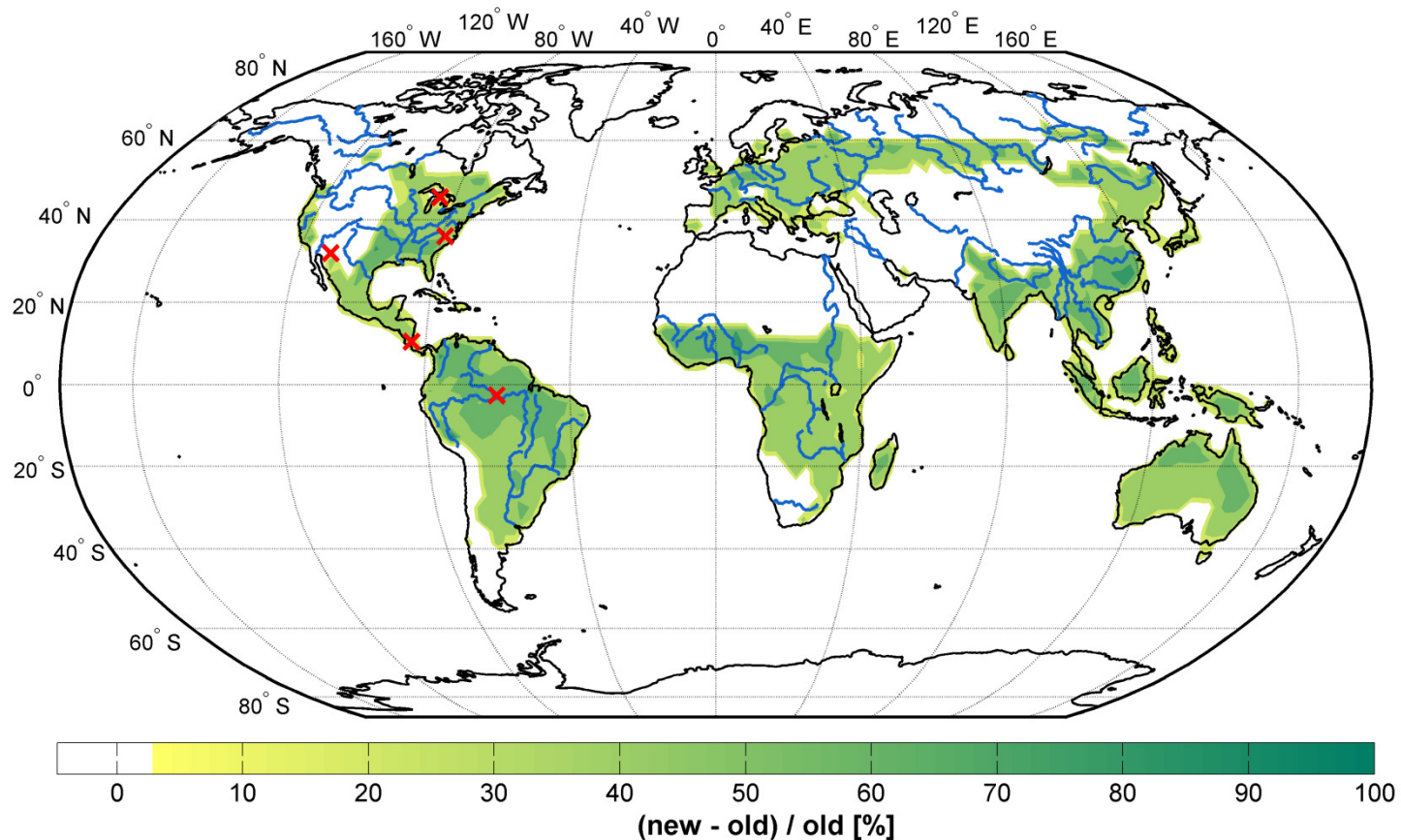
**THANK YOU**

2011

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**ADDITIONAL SLIDES**



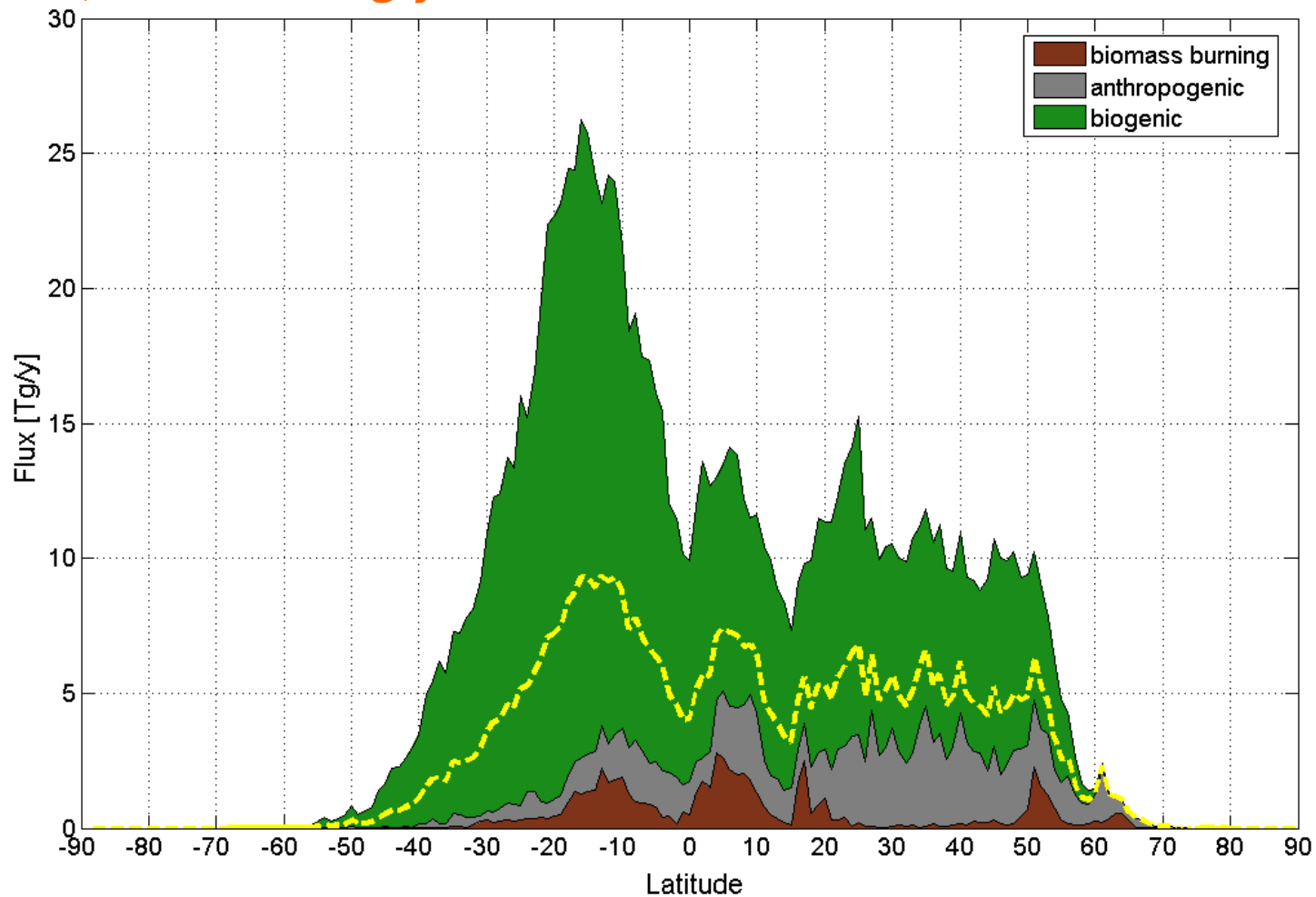
# Deposition

	Mean [TgC/a]	Comments
<b>this study</b>	<b>590±130</b>	<b>Dry and wet deposition (vapors)</b>
Goldstein and Galbally (2007)	200±100	Dry and wet deposition (vapors)
Hallquist et al., 2009	800	Dry and wet deposition (vapors)
Willey et al. (2000)	430±150	Wet deposition (vapors+particles)

# Latitudinal Distribution of NMVOCs

**Bottom-up (total): 1580 Tg/y**  
**IPCC2001, 2007: 450 Tg/y**

Compare to CH<sub>4</sub>: about 500-600 Tg/y  
Isoprene: about 500-700 Tg/y



# Biogenically enhanced secondary organic aerosol formation?



*Blue Haze above the OZARKS*

Some studies suggest (e.g. Jacobson, JGR, 2010) that a reduction of methane and black carbon aerosol might be an effective short-term climate mitigation strategy. This would be an example of a *win – win situation* (air quality and climate benefit).



# AQ-climate interaction

- Changes in  $\text{NO}_y$  (*loose?-win?*)
  - increases in reactive nitrogen leads to increases in ozone (*climate penalty/AQ penalty*) and increases in the oxidation capacity of the atmosphere (via primary OH production and  $\text{NO}_x$  recycling) (*climate benefit/AQ impact ? – (AQ penalty if it leads to more SOA)*) – this in turn would lead to a decrease in atmospheric lifetimes of reactive climate active gases (methane, HFC) (*climate benefit/AQ benefit*)
  - increase in scattering aerosols (enhanced SOA formation) (*climate benefit/AQ penalty*)

## Reduction of aerosols

Aerosols are predominantly thought to exhibit a negative forcing on climate (exception black soot – positive forcing)

*Win-loose*: cleaning the air from aerosols can have a climate penalty, but result in an air quality benefit

*Win-Win (black soot, methane)*: reduction in black carbon and methane (air quality and climate benefit).